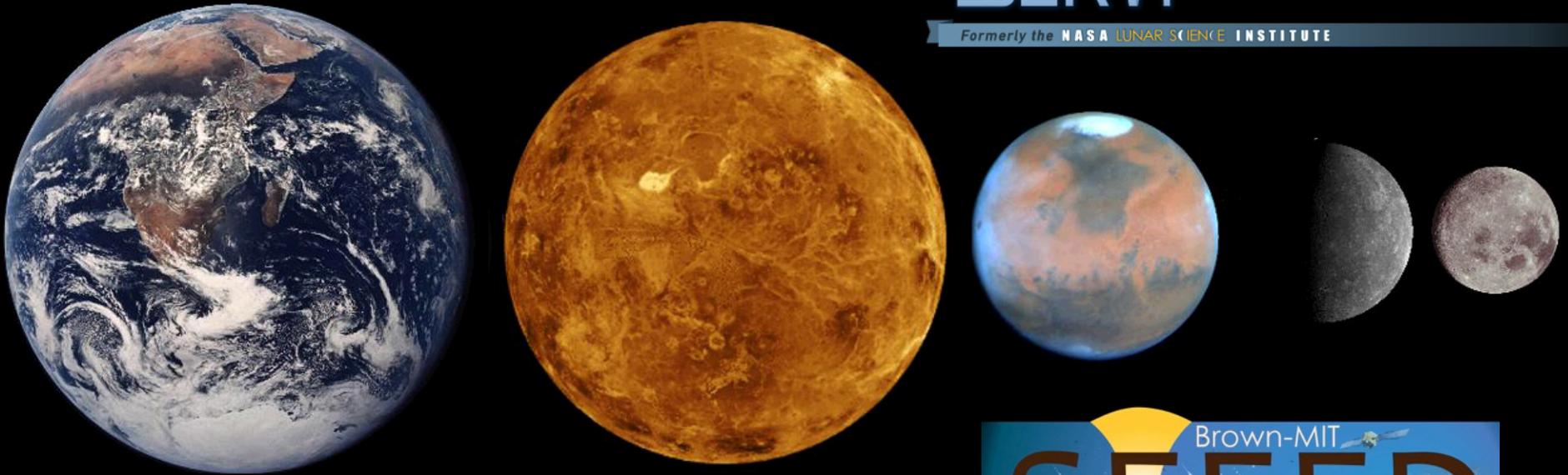


Ascent and Eruption of Magma on the Terrestrial Planets: A Planetary Perspective on the Moon

SERVI

SOLAR SYSTEM EXPLORATION RESEARCH
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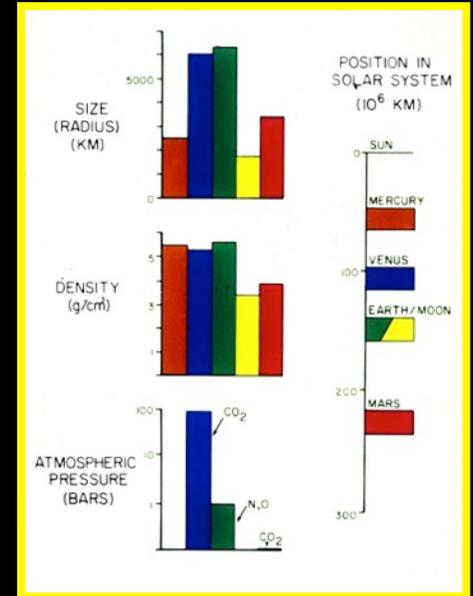
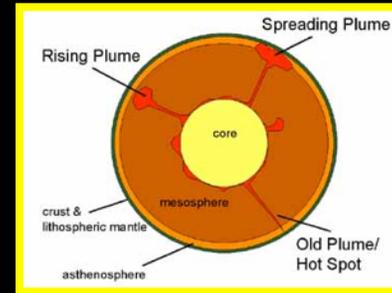


James W. Head¹ and Lionel Wilson².

¹Department of Earth, Environmental and Planetary Sciences,
Brown University, Providence, RI 02912 USA,

²Lancaster Environment Centre, Lancaster University, Lancaster LA1 4YQ, U.K.

Generation, Ascent, and Eruption of Magma:



- Analysis on Earth and Terrestrial Planets:
 - Provides substantial information about the geological history/thermal evolution of each body.
- Synthesis of Array of Extrusive Features & Landforms:
 - Provides insight into eruption styles, lithospheric stress states, and mantle convection: A Grand Experiment!
- Surface of Terrestrial Planet Elemental Compositions:
 - Consistent with a range of mantle compositions, but all are likely to produce mafic to ultramafic melts.

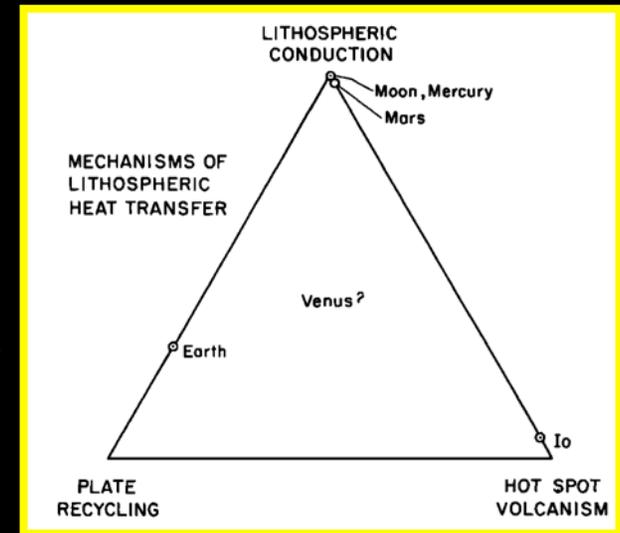
Main Controls on the Types of Surface Volcanic Features and Accumulations



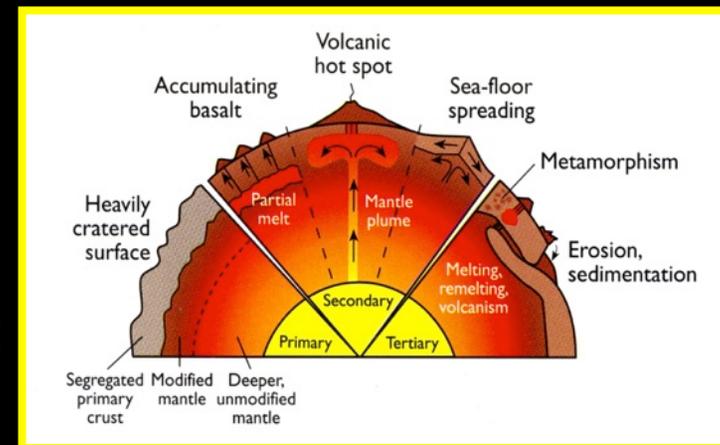
Differences in:

- 1) Magma composition/volatile content.
 - 2) Lithosphere continuity and thickness.
 - 3) Tectonic regimes.
 - 4) Crustal origins and stability.
 - 4) Crustal thicknesses and densities.
 - 5) Mantle thickness.
 - 6) Mantle convective style.
- Preferred locations for magma reservoirs:
- 1) at depth within a planetary interior or
 - 2) relatively shallow within a volcanic edifice.

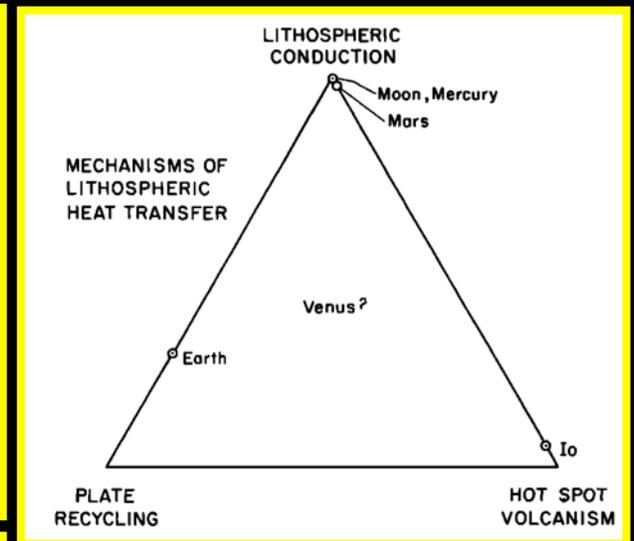
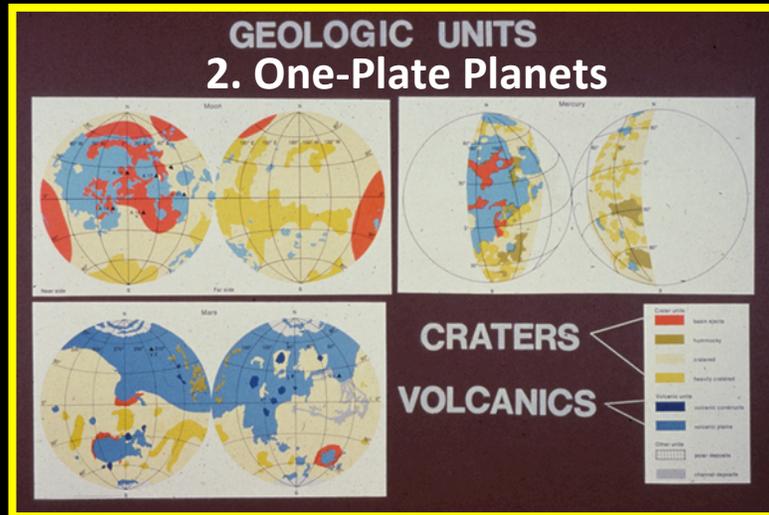
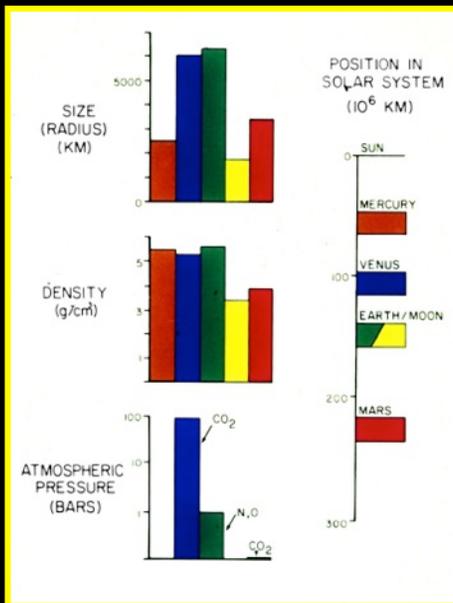
Modes of Planetary Heat Loss



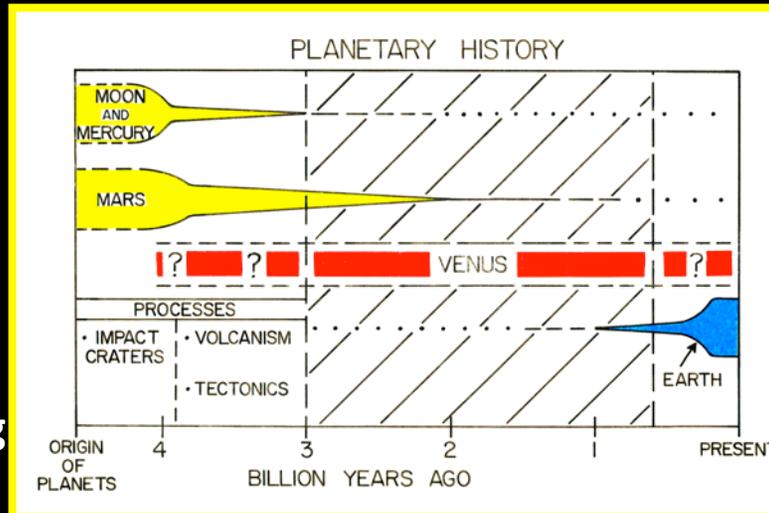
Modes of Crustal Formation



Planetary Heat Loss Mechanisms: Thermal Evolution



4. Planetary Resurfacing



- Global stress regimes characterized by net horizontal extension favor upward propagation of dikes, whereas horizontal compression inhibits dike propagation.

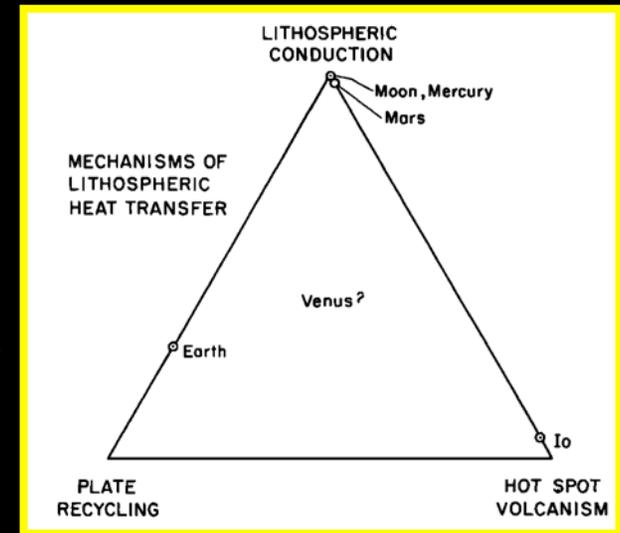
Main Controls on the Types of Surface Volcanic Features and Accumulations



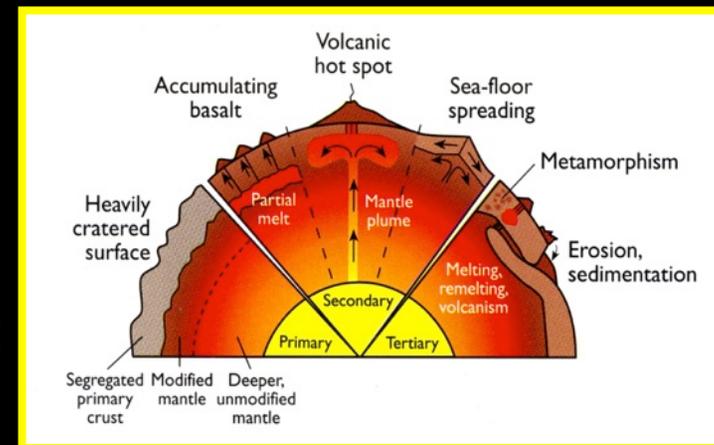
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Modes of Planetary Heat Loss



Modes of Crustal Formation

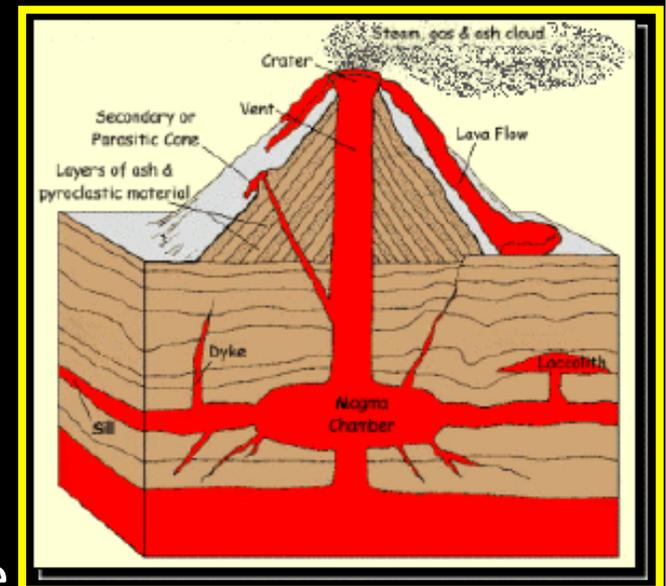
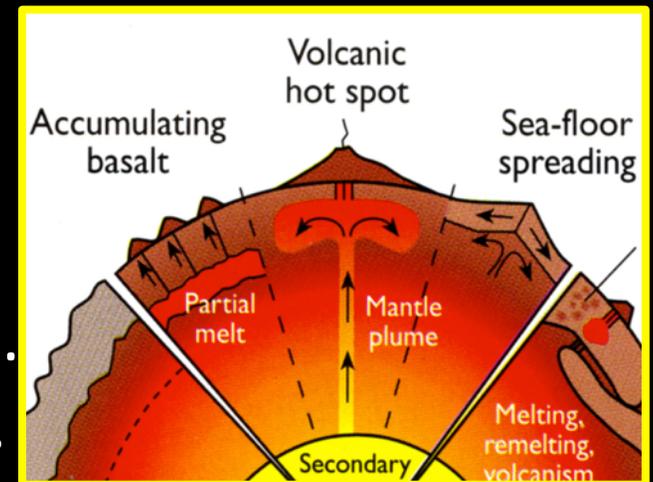


Main Controls on the Types of Surface Volcanic Features and Accumulations



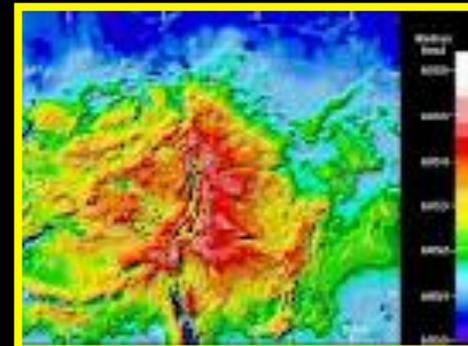
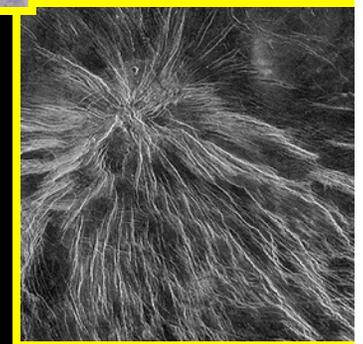
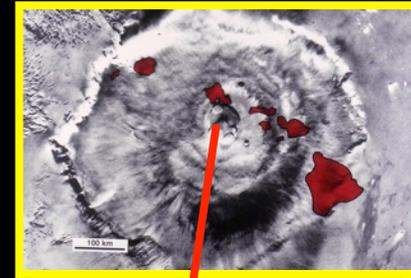
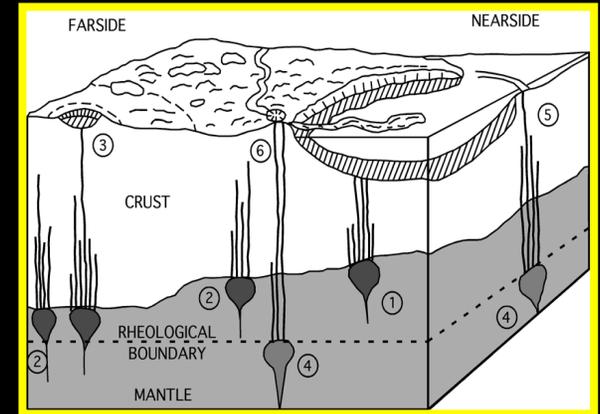
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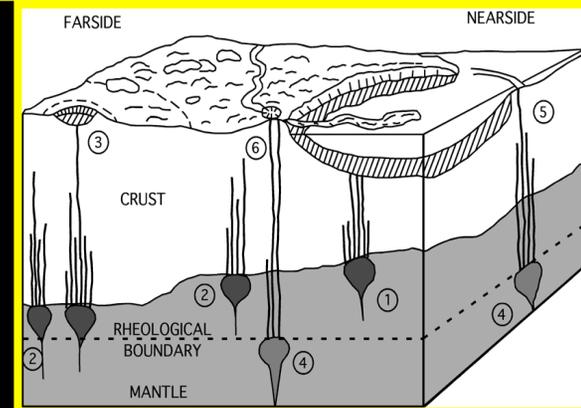
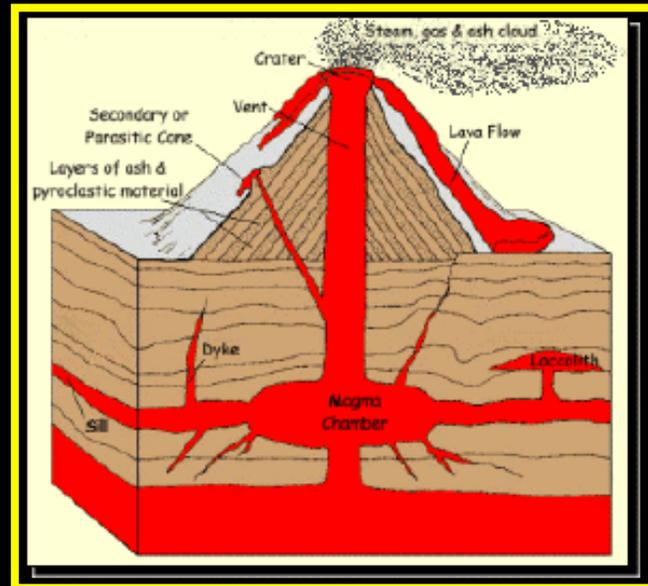
Magma Reservoir Formation

- **Deeper reservoirs:**
 - Can form near the **rheological change** at base of the lithosphere.
 - At upwellings, due to **pressure-release melting**, or
 - At vertical **discontinuities in density** such as at the base of the crust.
- Evidence for reservoirs in edifices is seen in **calderas**.
- **Floor-fractured craters** represent the surface manifestation of shallow reservoirs on the Moon.
- Evidence for **deeper magma bodies** is seen in giant dike swarms.
- The position of **ascending mantle flow** is marked by broad rises formed from thermal uplift, enhanced crustal construction, and individual edifices built by surface eruptions.

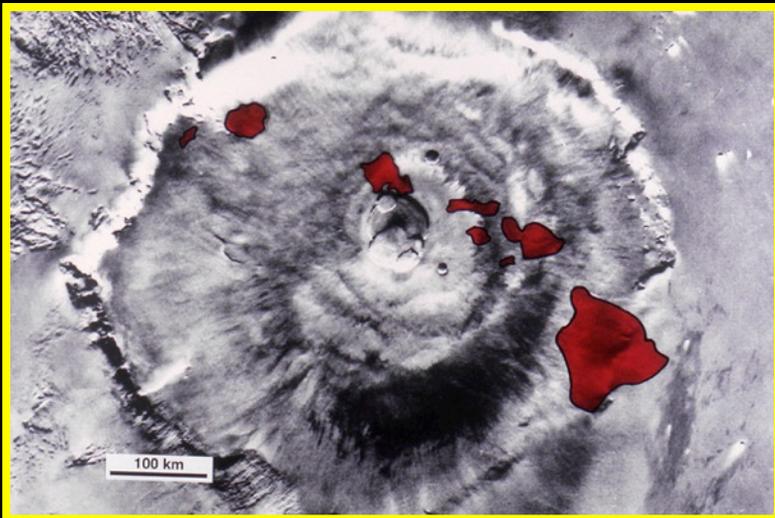


Shallow Magma Reservoirs

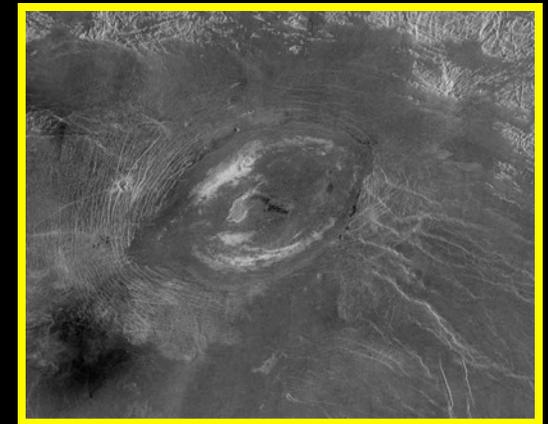
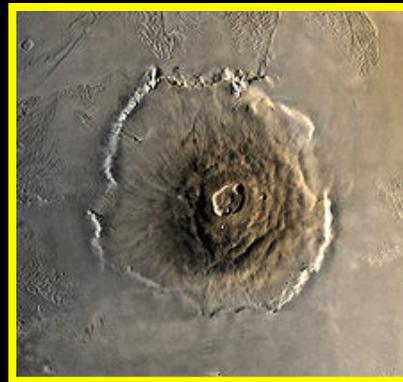
- Commonly formed within volcanic edifices on Earth, Mars, and Venus.
- Building a volcanic edifice and reservoir requires:
 - Multiple pulses of magma to rise frequently within a spatially restricted region over extended period of time.
- On the Moon, in contrast,
 - Low eruption frequencies and great flow lengths ensure that typical large edifices will not form.
- Shallow reservoirs form within edifices at levels of neutral buoyancy.
 - Repeated, relatively small-volume eruptions from the shallow reservoirs progressively build shield volcanoes of a range of sizes and aspect ratios on Earth, Mars, and Venus.
- **Calderas:** Large shield volcanoes commonly host summit collapse calderas, produced when substantial volumes of magma are erupted on the volcano flanks.



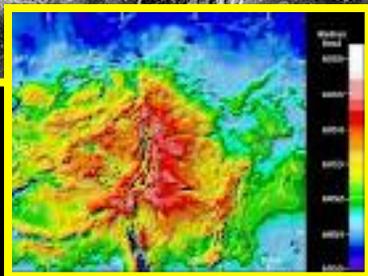
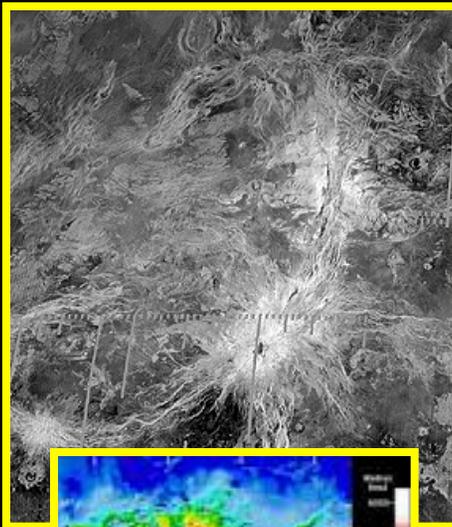
Typical Volcanic Features on the Terrestrial Planets



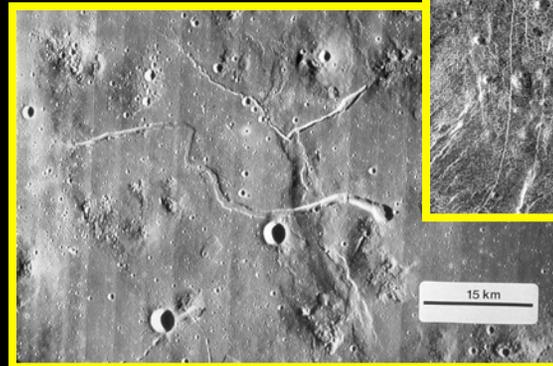
Major Shield Volcanoes.



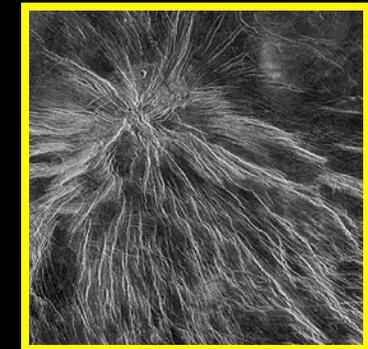
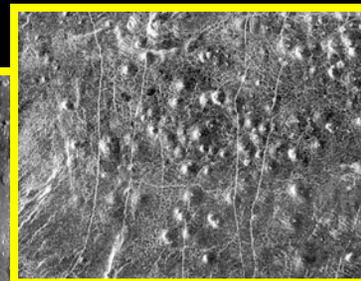
Calderas.



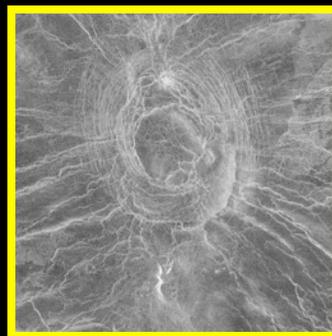
Rifted Rises.



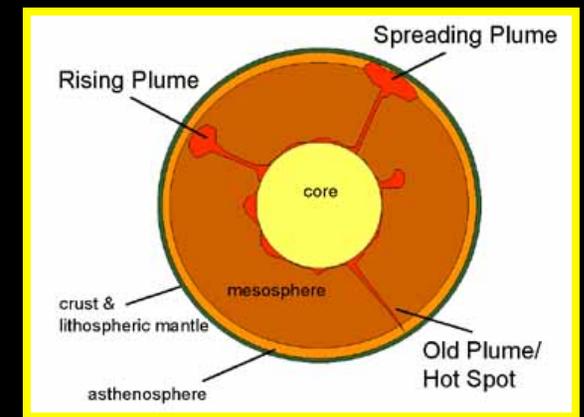
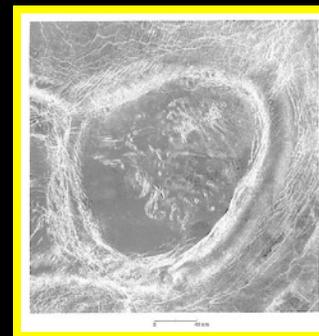
Shield Fields



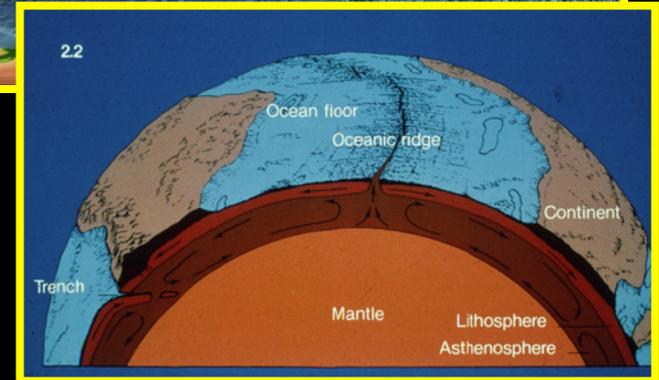
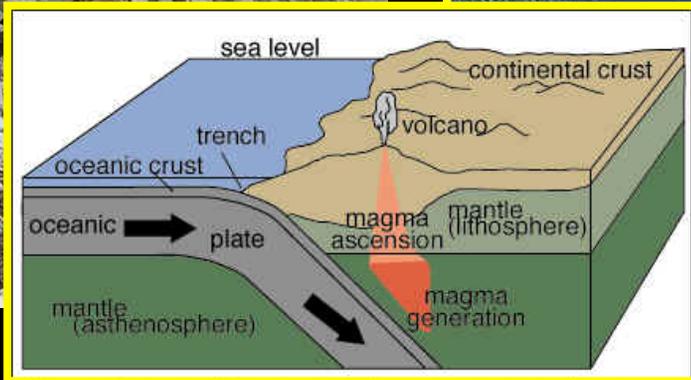
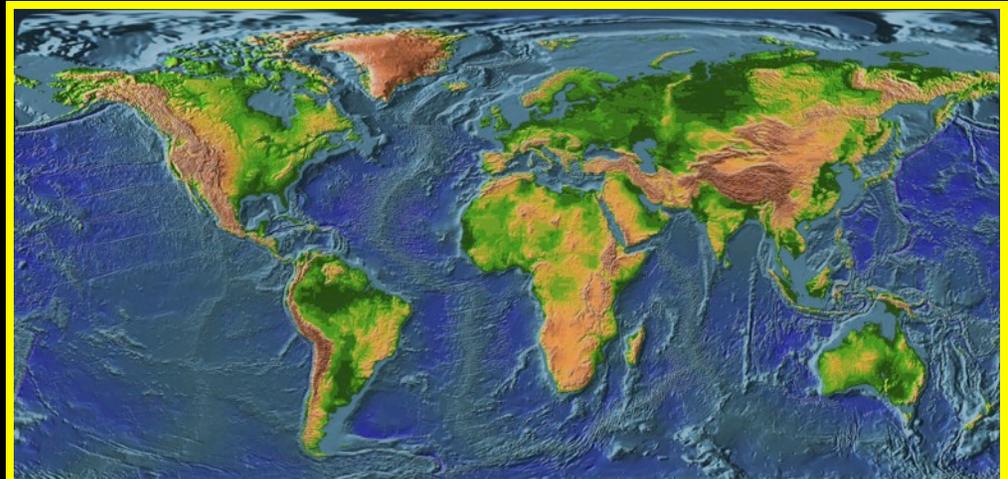
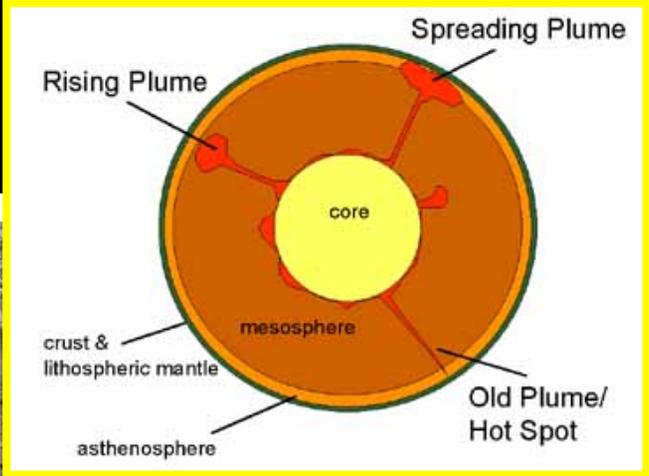
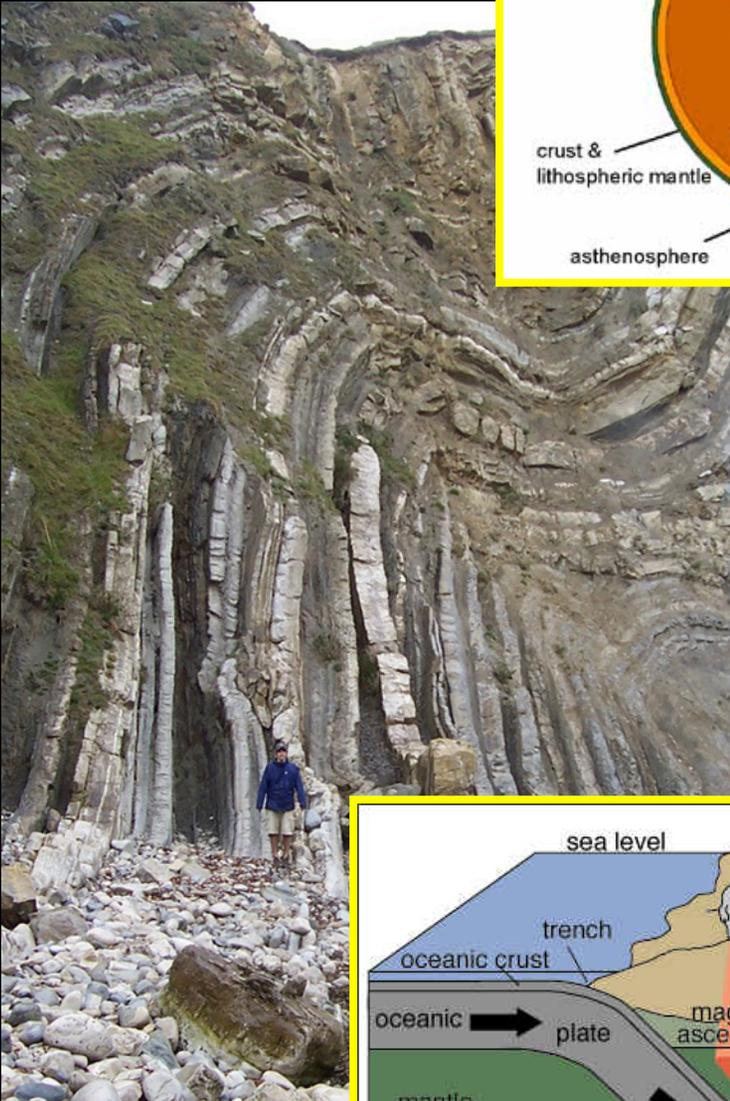
Novae.



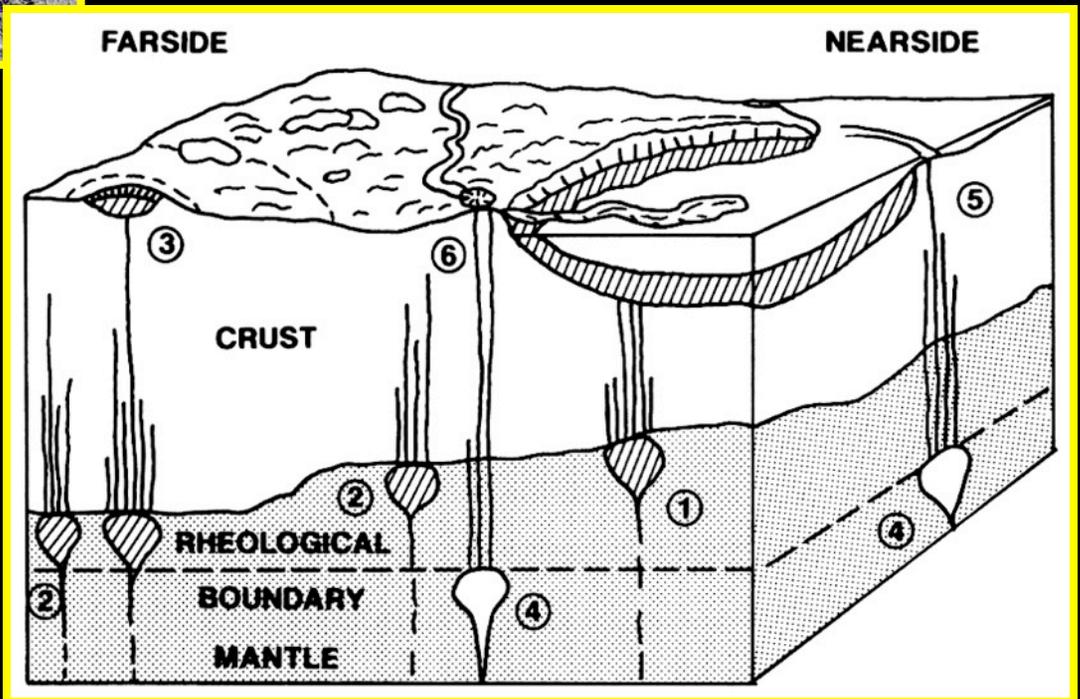
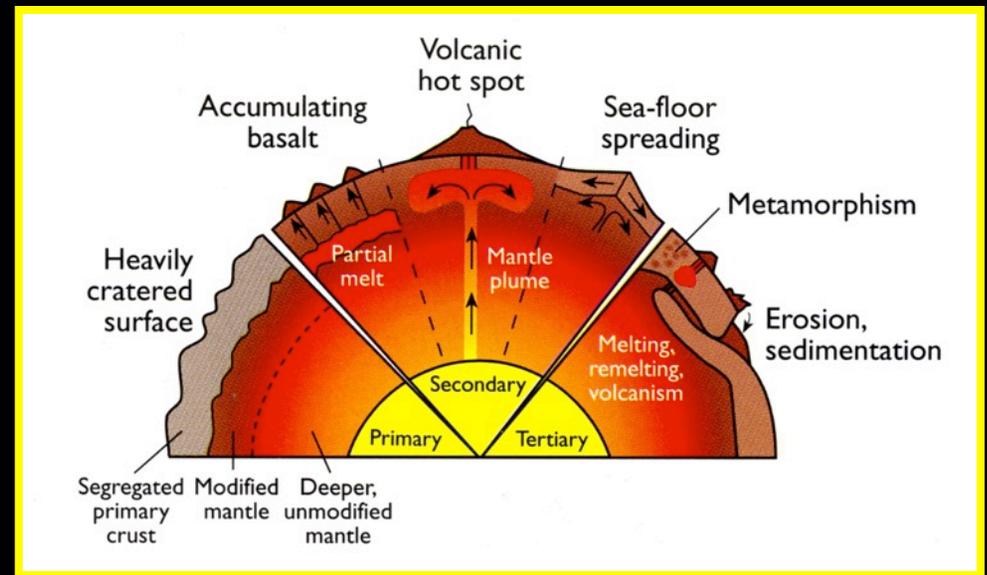
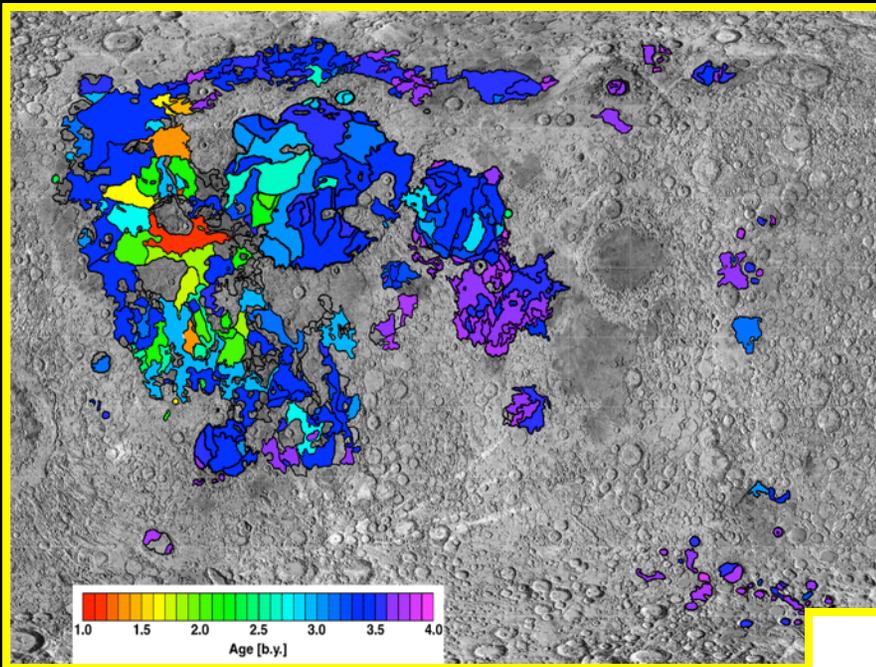
Coronae.



Earth



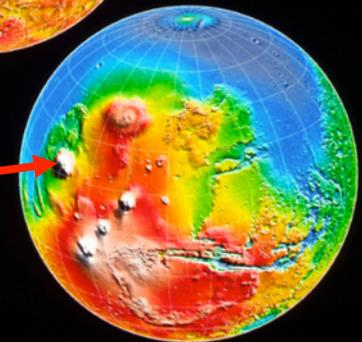
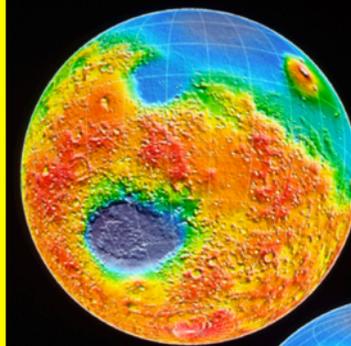
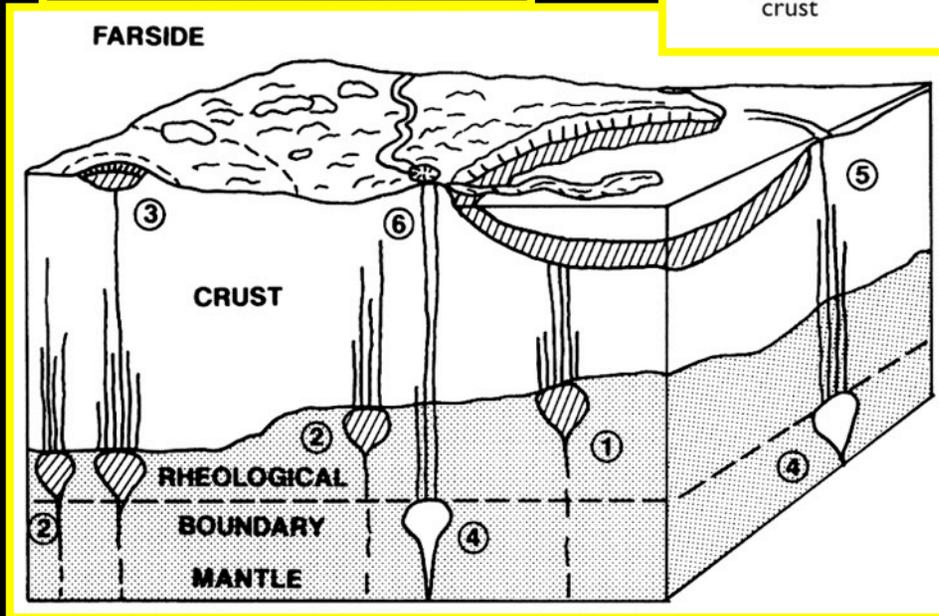
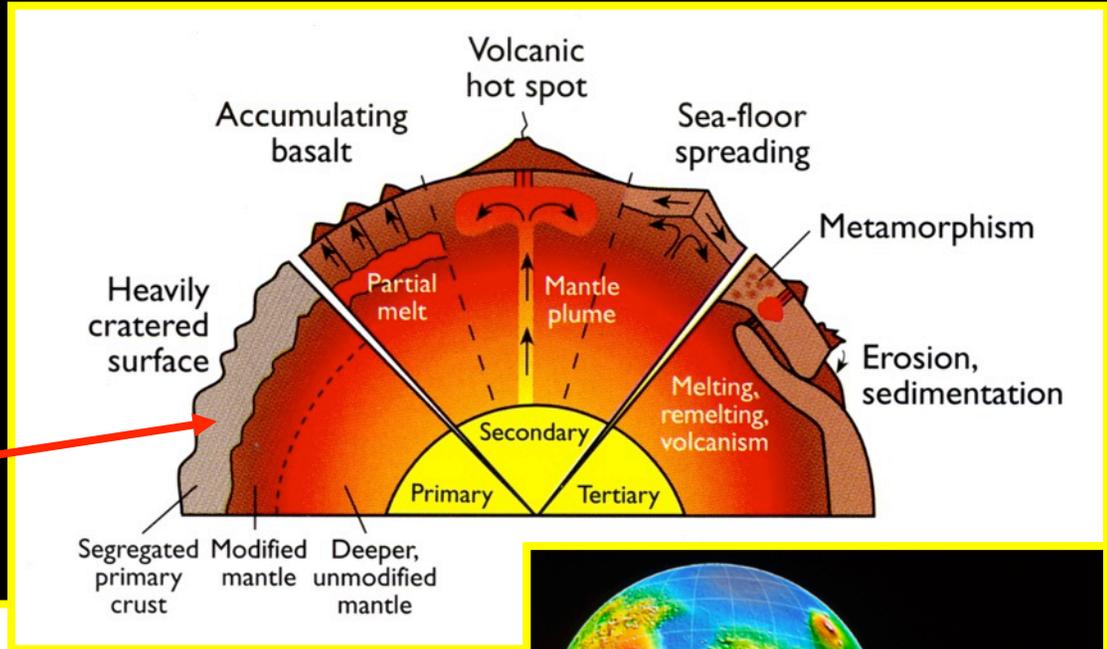
Planetary Crustal Formation and Lunar Volcanism



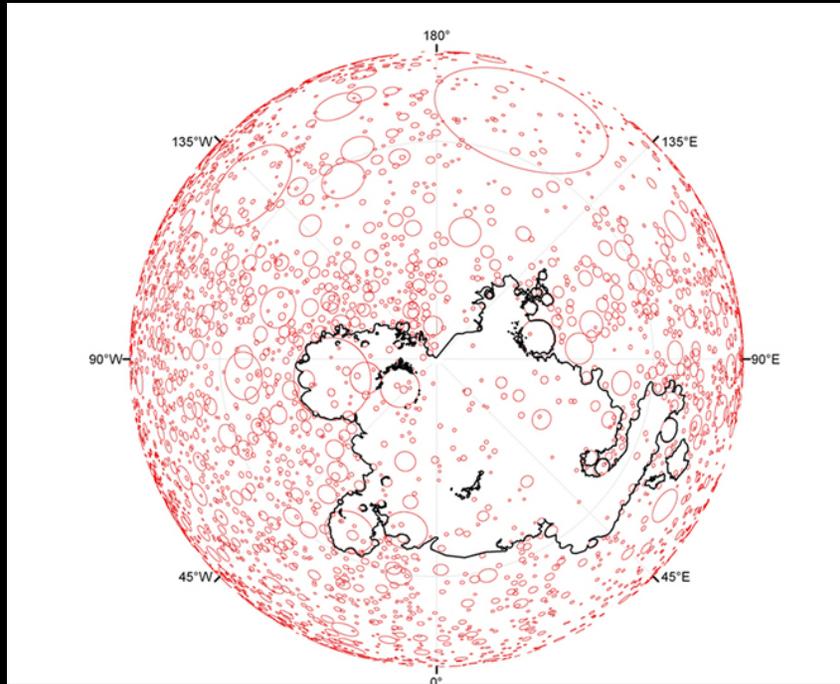
Some Basic Points:

1. Primary crustal formation.
2. Secondary Crustal Formation- Mare Volcanism.
3. Neutral buoyancy zones.
4. Cooling, shrinkage and inhibition to eruption.
5. Very high effusion rates/volumes.
6. Nearside-farside difference.

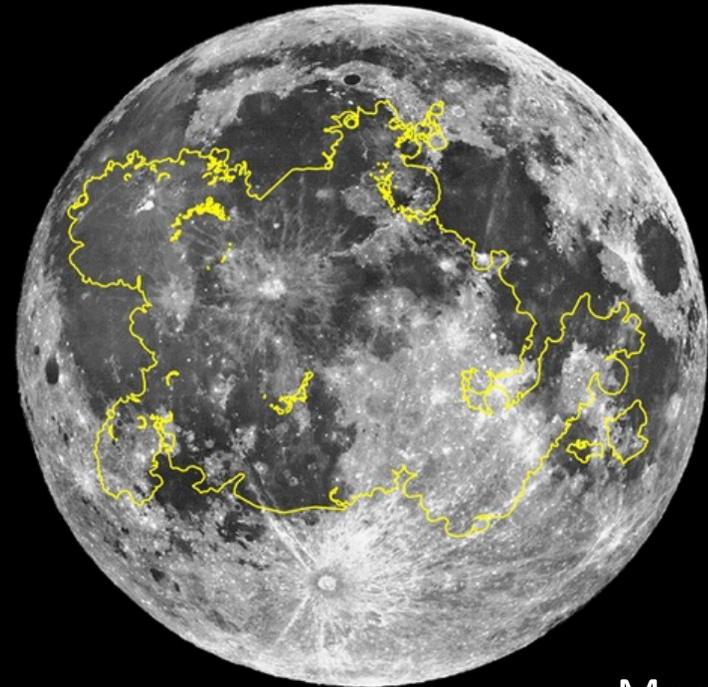
The Moon: Primary Anorthositic Crust, One Plate Planet, Cools Rapidly, Deep Magma Sources



Mercury: Areal Extent of the Northern Plains - A Comparison



Mercury



Moon

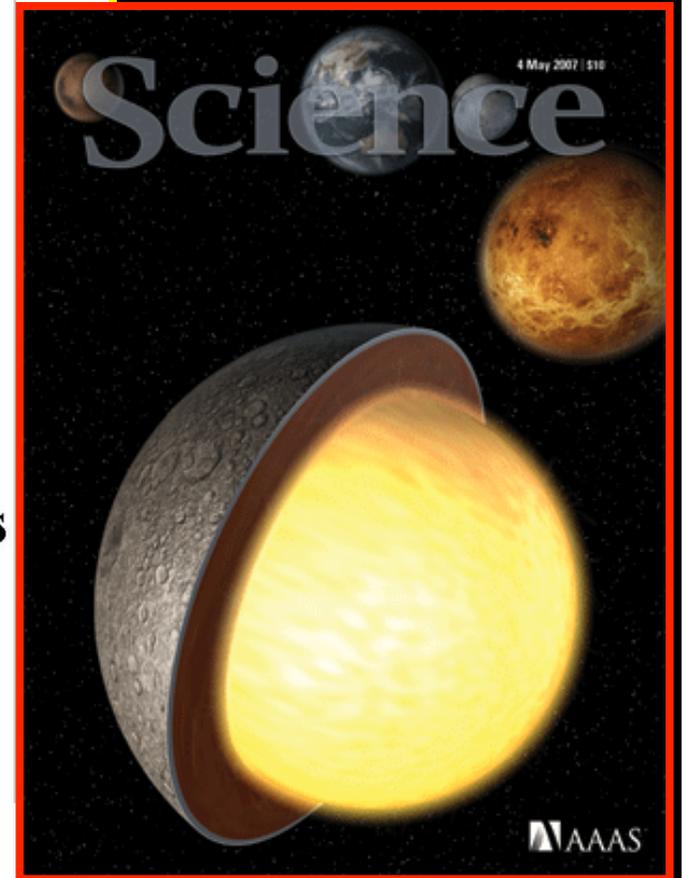
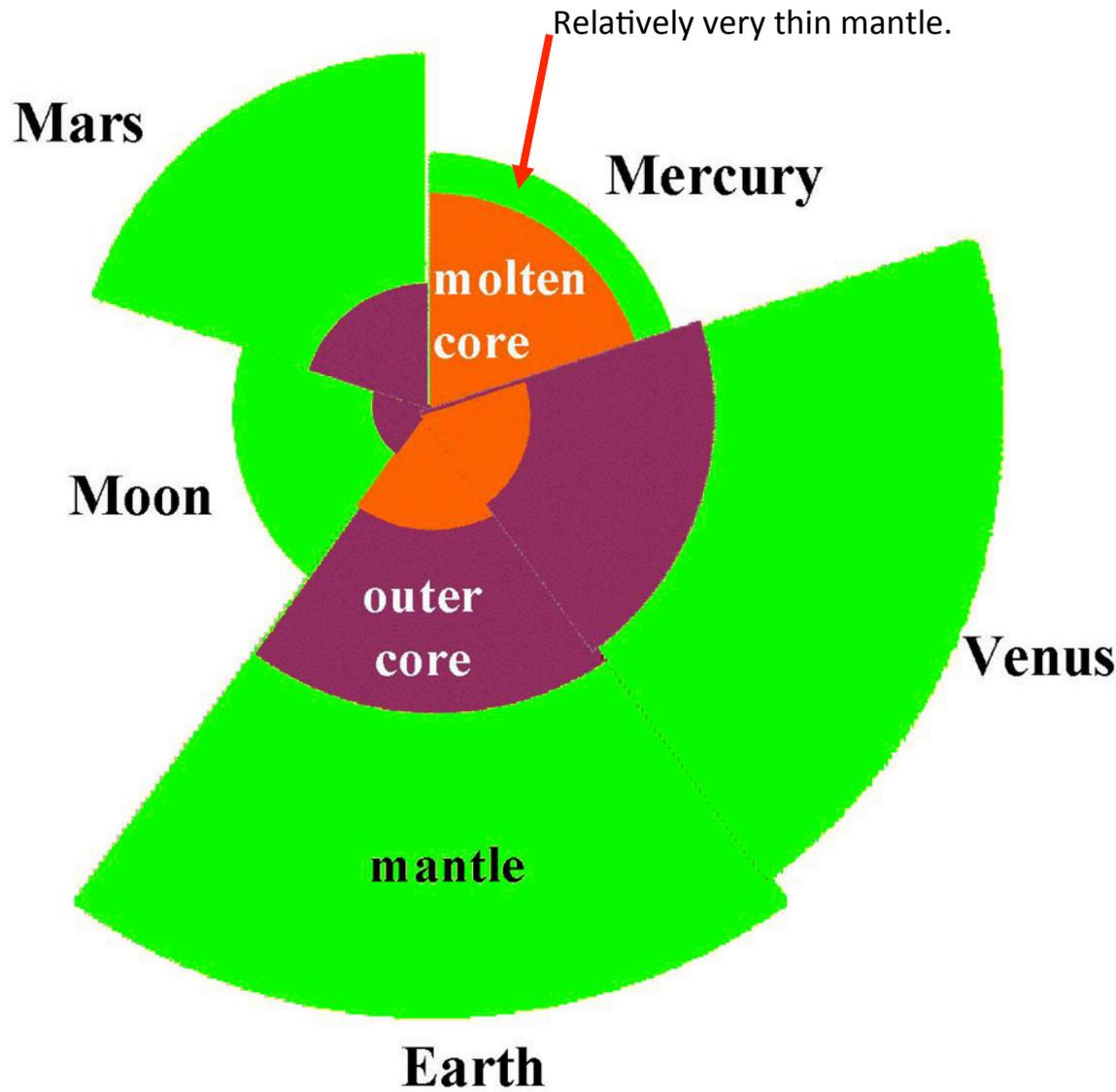


Earth

What is the nature and mode of emplacement of the northern plains?

(Head et al., 2011)

How Does the Mantle of Mercury Compare to Other Terrestrial Planets?



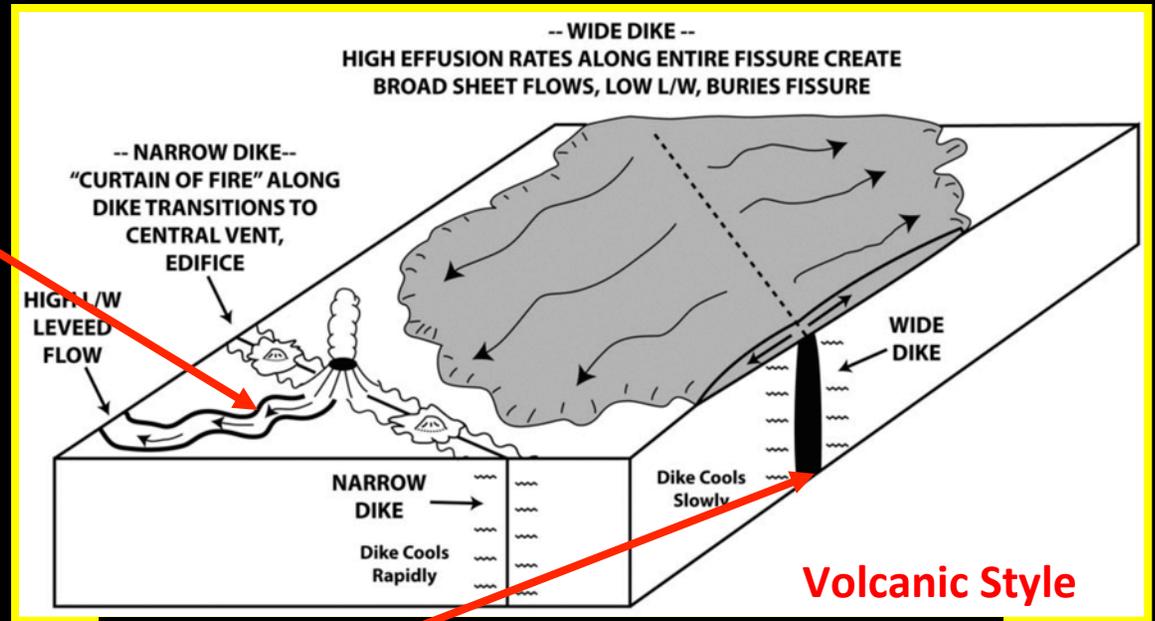
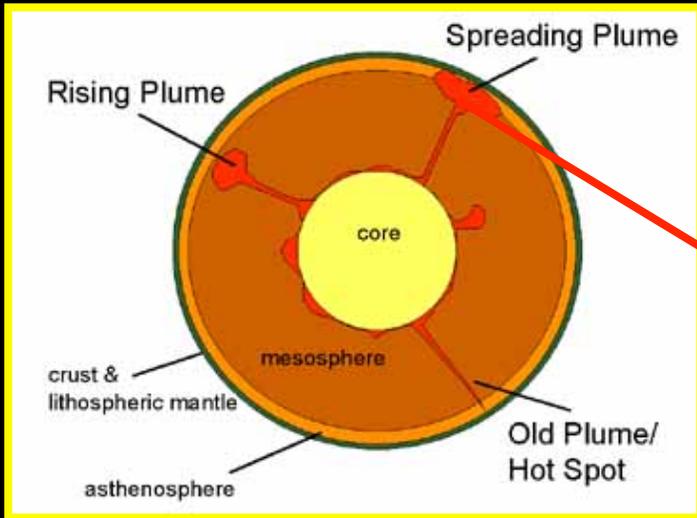
(G. Rieke, U. of A.)

(C. Bickel, *Science*)

Style of Mantle Convection, Heat Loss, Volcanism

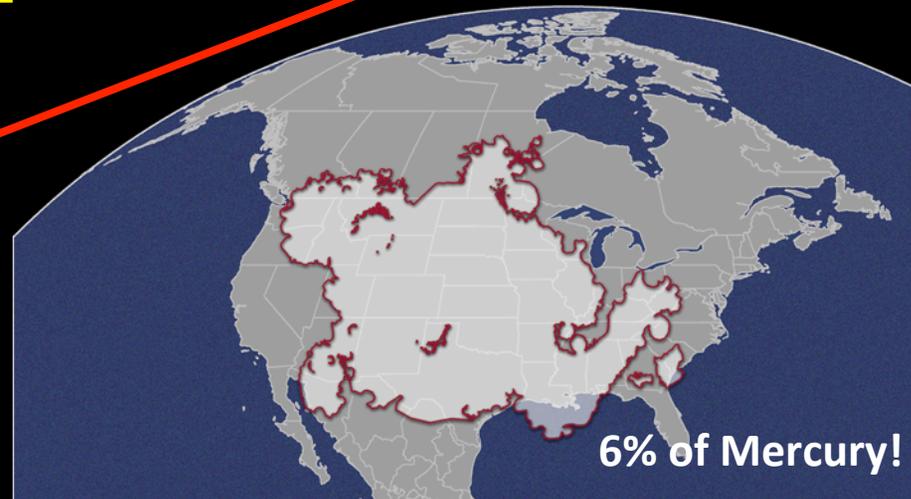
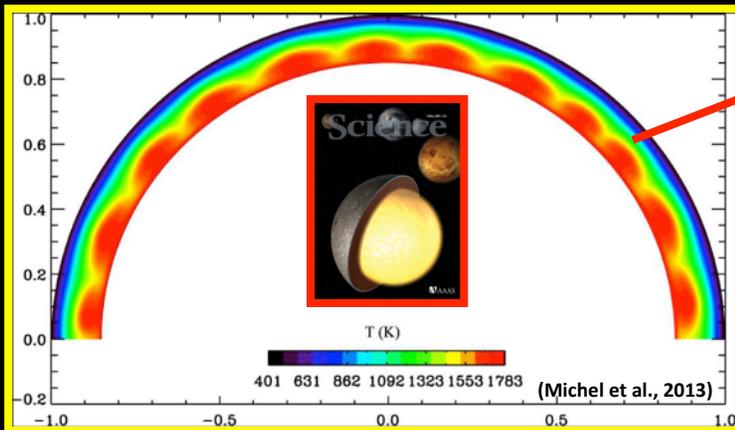
(Two End-Members)

1. Thick Mantle: Convecting Plumes.

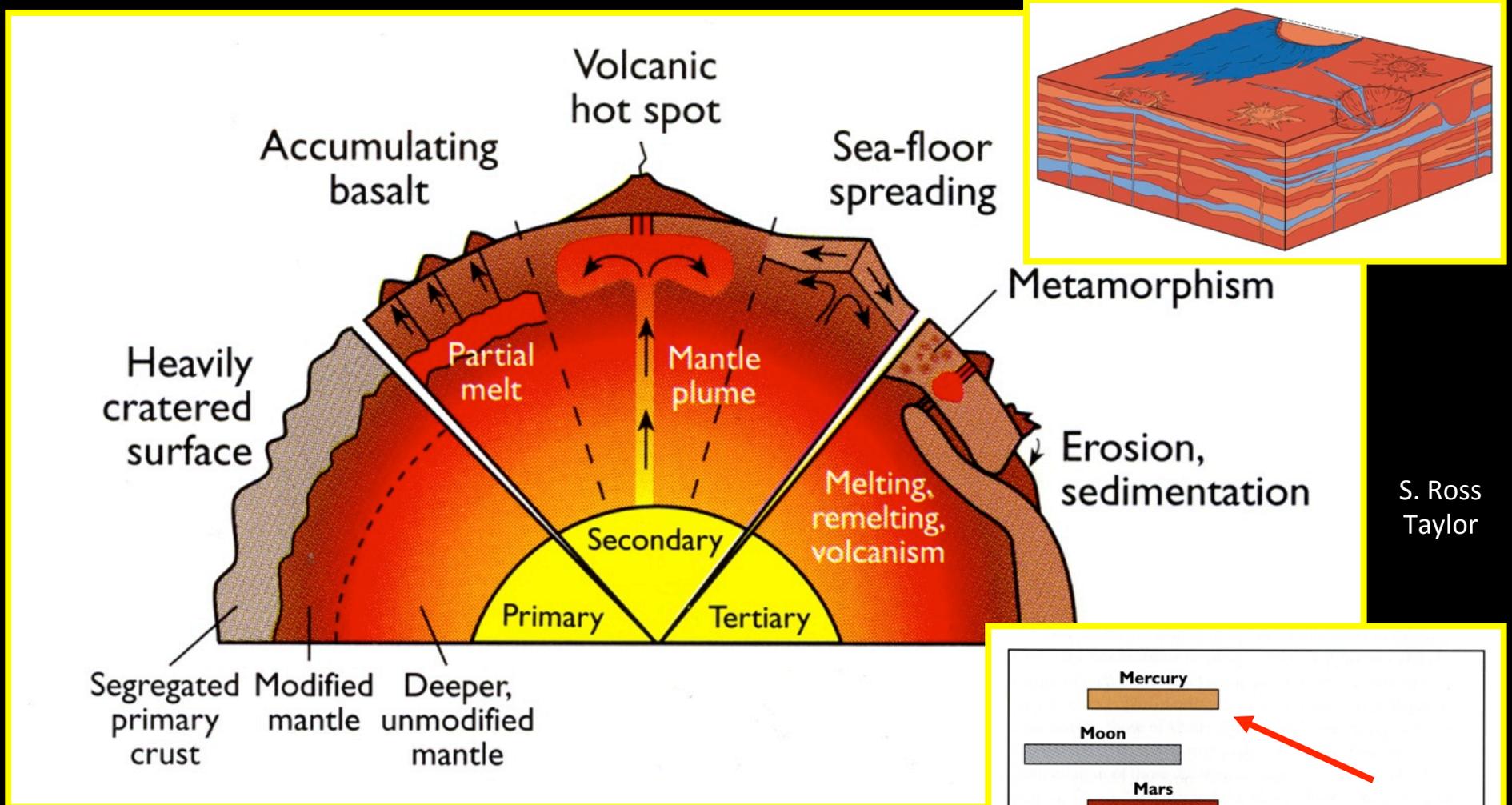


Volcanic Style

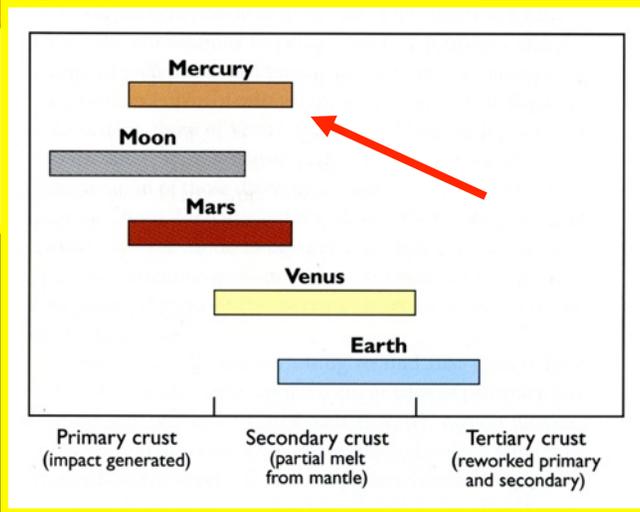
2. Thin Mantle: Convection Inhibited.



Crustal Formation and Evolution: Primary, Secondary, Tertiary

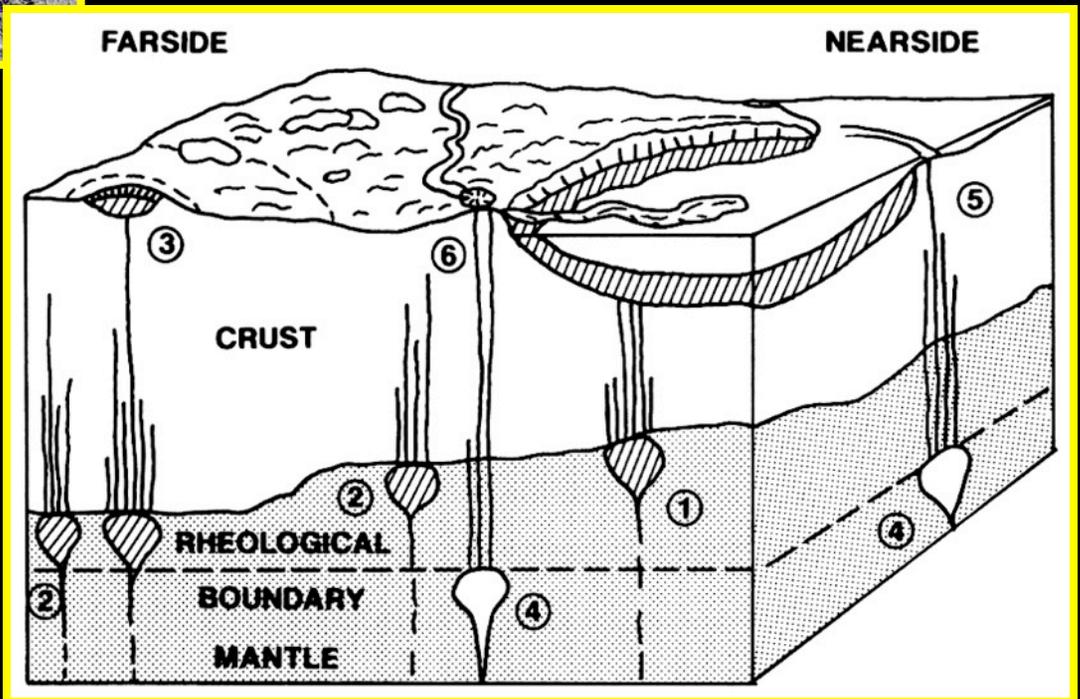
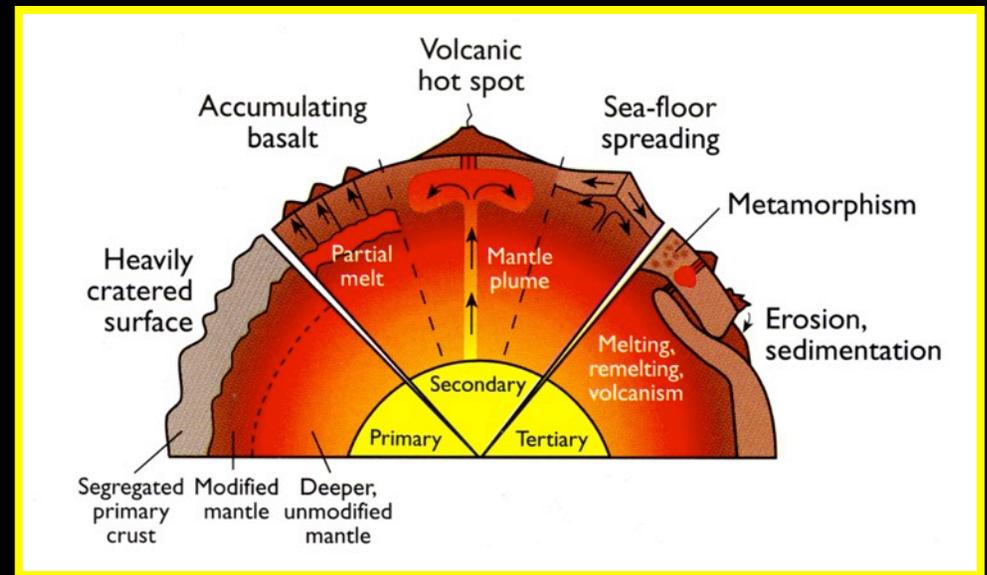
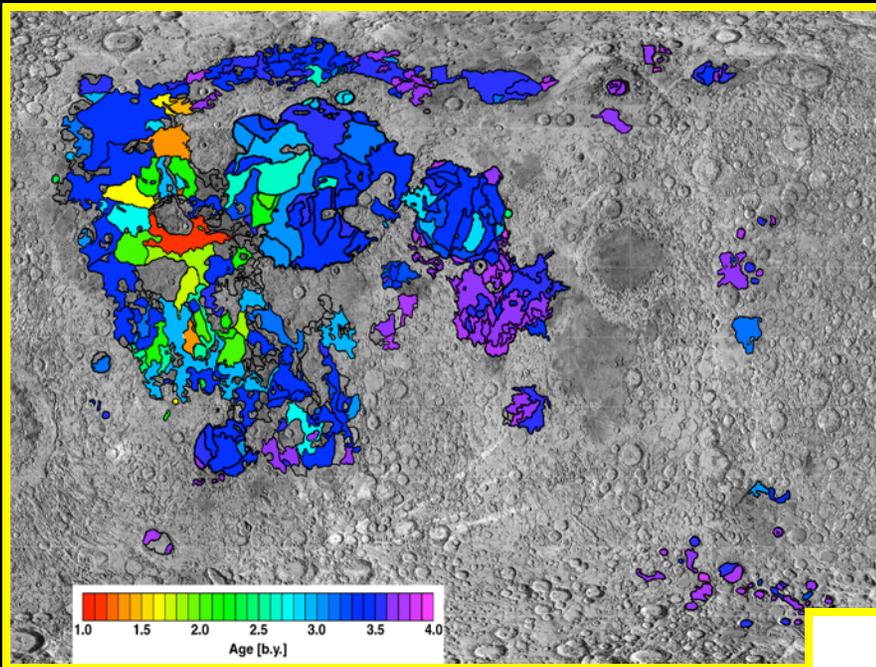


S. Ross Taylor



- No positive evidence for a Moon-like plagioclase flotation primary crust.
- Abundant evidence for global, thick volcanic deposits: Similar to secondary crust on the Moon and Mars.
- What is the nature of primary crust on planets like Mercury?

Planetary Crustal Formation and Lunar Volcanism



Some Basic Points:

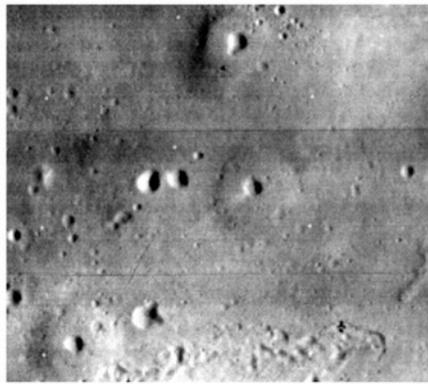
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Range of Lunar Volcanic Landforms

How can this diversity be explained?



Cones



Small Domes



Pits



Crater Chains



Linear Rilles



Floor-Fractured Craters

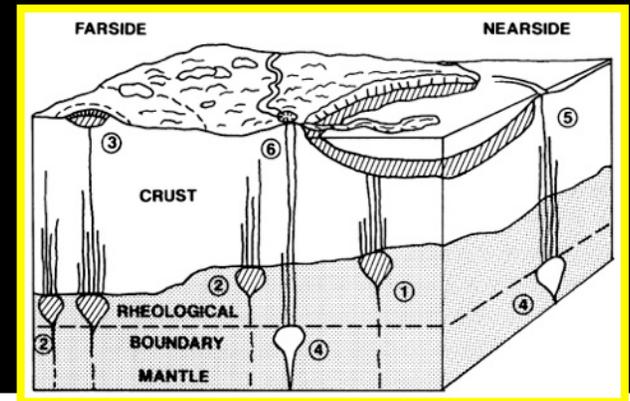
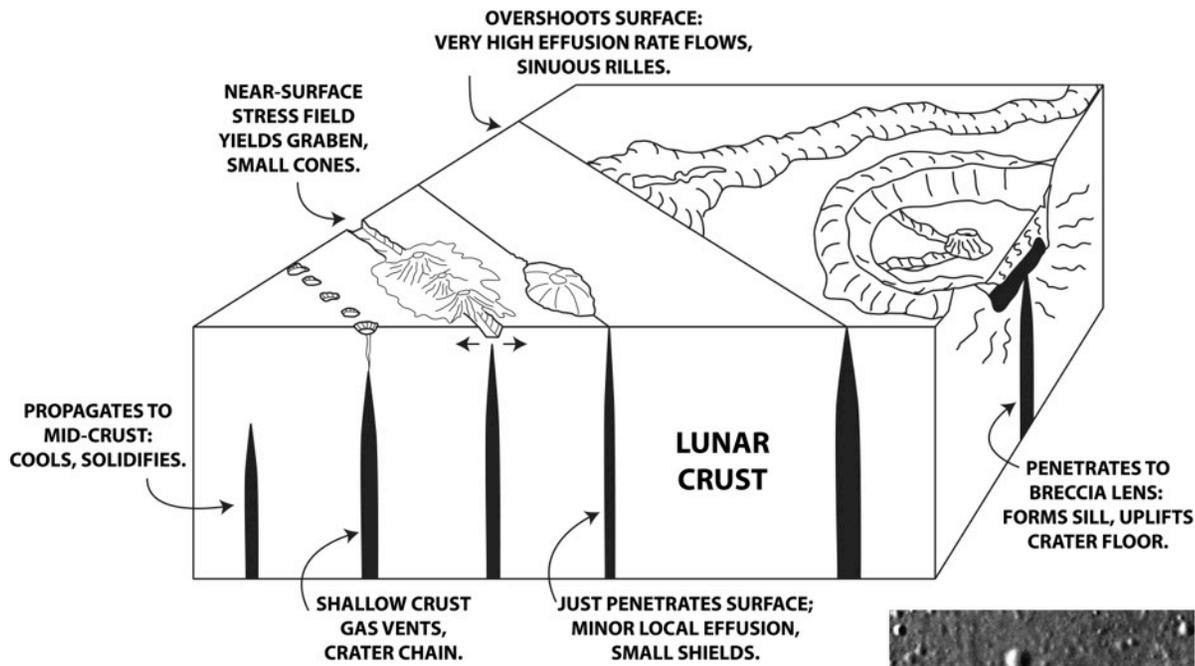


Mare Imbrium

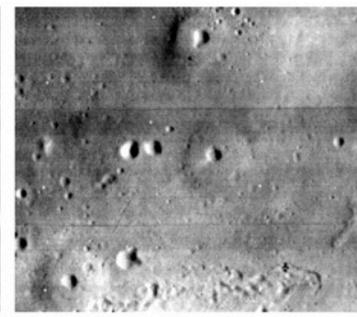


Rimae Prinz

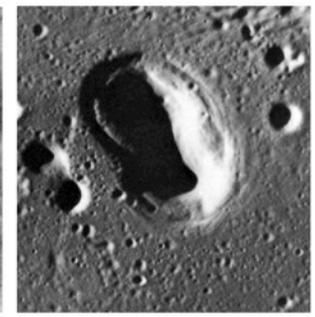
Spectrum of Overpressurization and Dike Emplacement Events



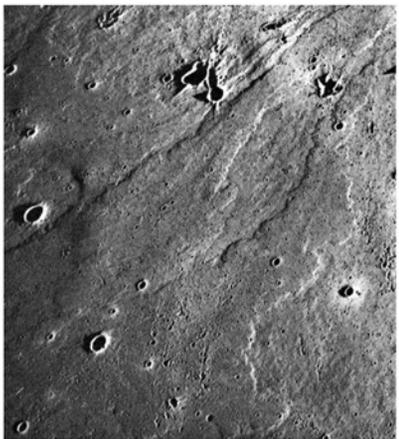
Cones



Small Domes



Pits



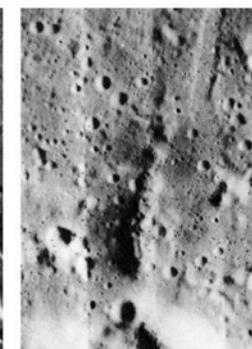
Mare Imbrium



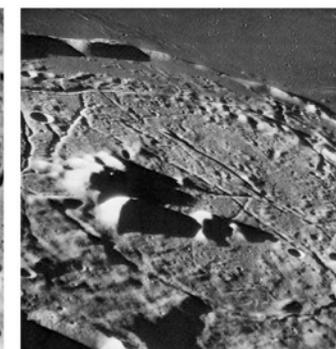
Rimae Prinz



Crater Chains



Linear Rilles



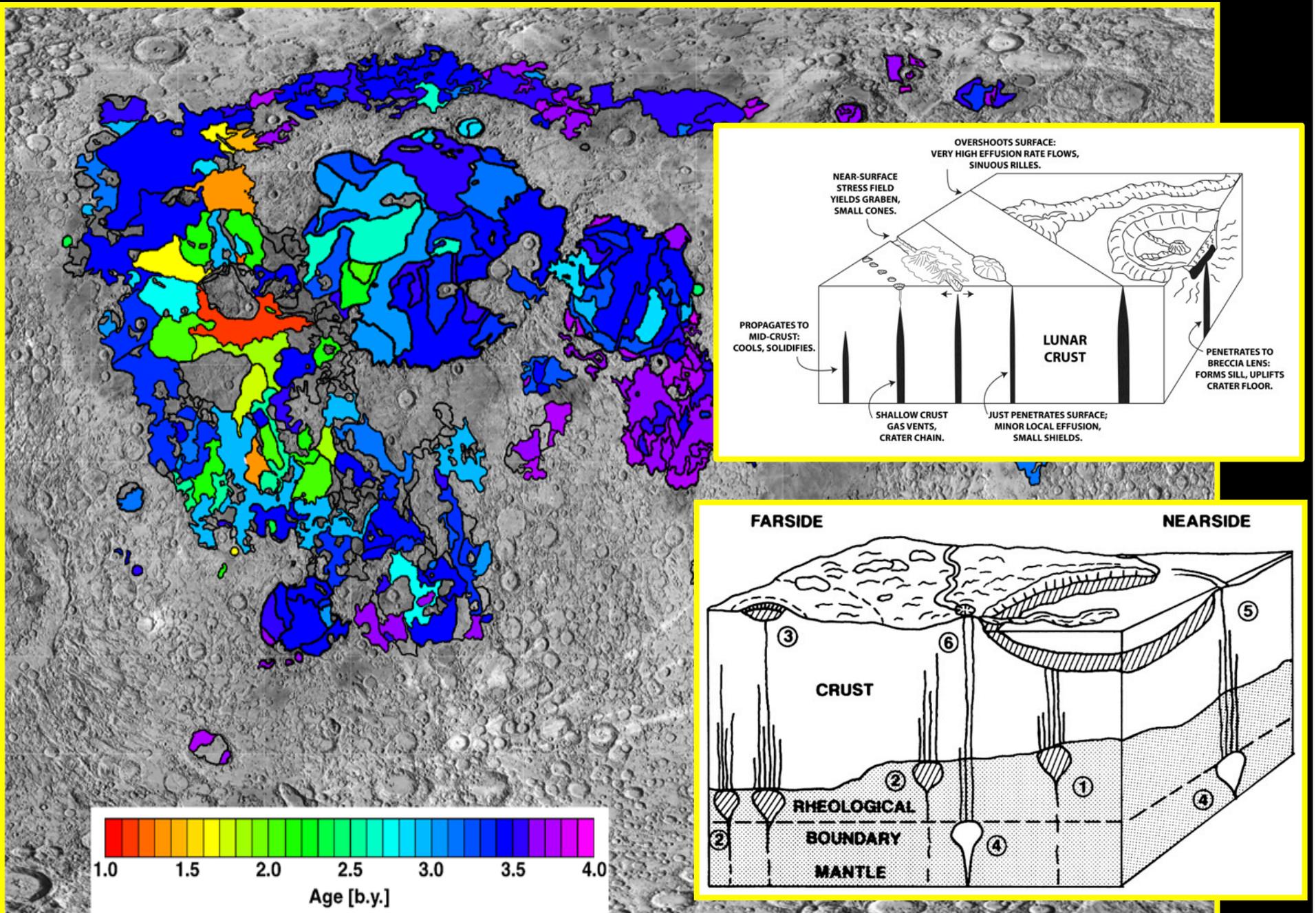
Floor-Fractured Craters



Future Human and Robotic Exploration of the Moon:

- What are the most likely locales for deep mantle xenoliths.
- What is the nature and origin of young mare basalt volcanism?
- What is the chronology of young mare basalt volcanism?
- What is the origin of Ina “caldera” and similar features.
- What is the relationship of lunar picrites, glasses and lava flows?
- Nearside/Farside mare basalt differences: Ages, mineralogy, volumes and processes.
- Human Exploration:
 - Geometry of dikes to test ascent and eruption models.
 - Magnetic properties, diameter, density, gravity.
 - Floor-fractured craters: Intrusion, alteration, hydrothermal activity.
 - Marius Hills: Gravity structure and relation to solidified magma reservoirs.
 - Sinuous rille source region:
 - Abundance of deep dikes in the lunar crust.
 - Emplacement dynamics of lunar pyroclastics: Vent structure and 3D characteristics. Role of volatile sources and sinks.
 - Dikes as recorders of the history of the magnetic field.
 - Cryptomaria: Are the GRAIL anomalies actually cryptomaria?
 - What is the role of impact basins in the generation, ascent, eruption and collection of magma?
 - How are spinel minerals related to dike emplacement.

Lessons for Human and Robotic Exploration of the Moon:



Future Human and Robotic Exploration of the Moon:

- Geological Exploration: Major Processes:
- Mineralogical Diversity and Origin:
- Petrological Diversity and Petrogenesis:
- Geophysics: Seismology, Gravity, Magnetics, GPR:
- Regolith Formation and Evolution:
- In Situ Resource Utilization:

