

Evolution of Gas and Foam in Lunar Floor-Fractured Craters

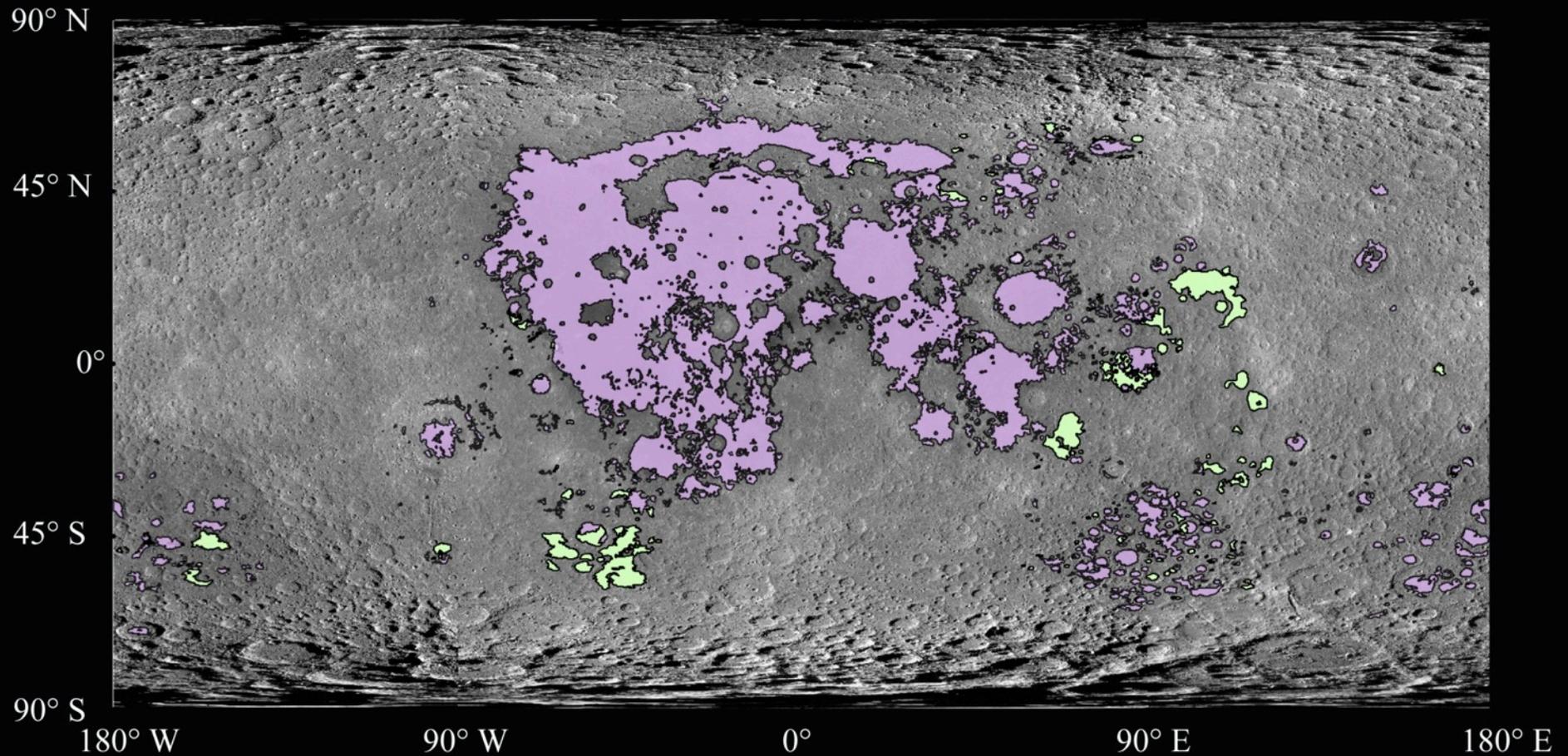
L. M. Jozwiak, J. W. Head, and L. Wilson

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SSERVI Exploration Science Forum



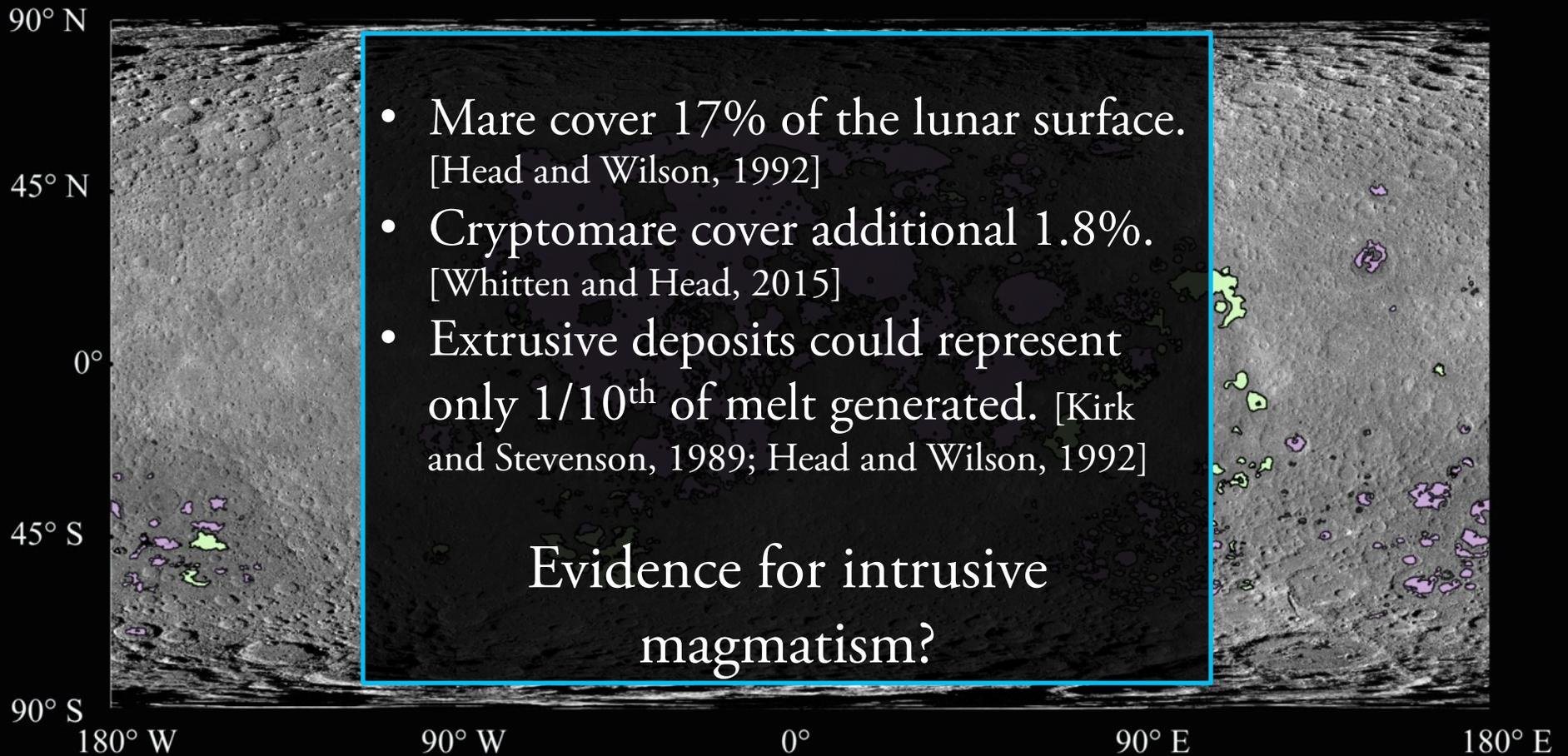
Lunar Volcanism



Mare Distribution: USGS

Cryptomare Distribution: Whitten and Head (2015)

Lunar Volcanism

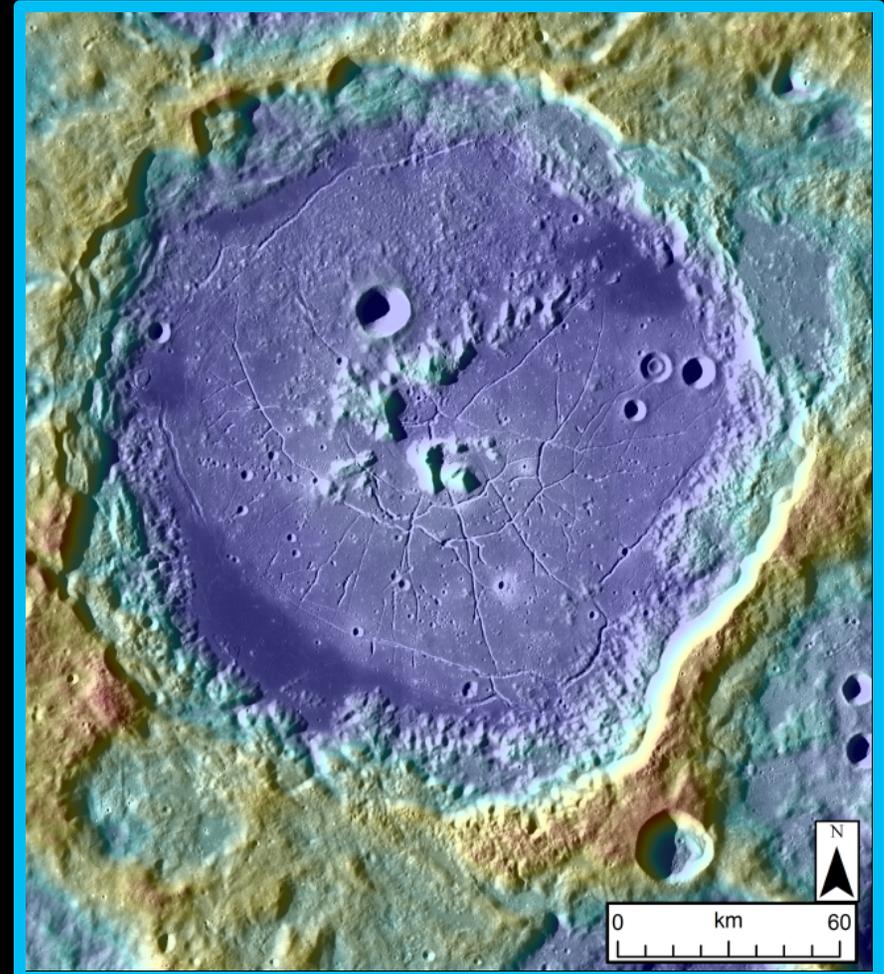


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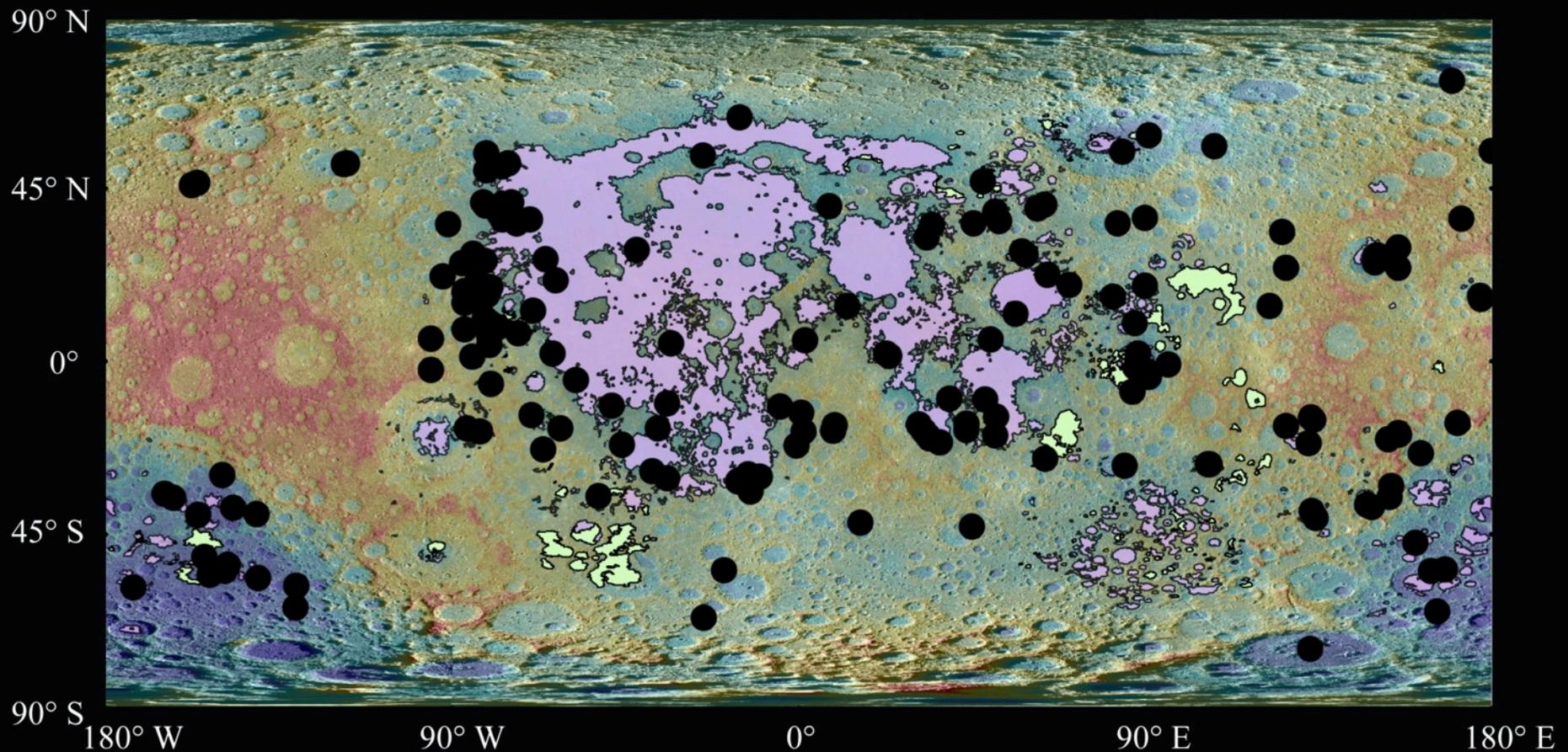
Lunar Floor-Fractured Craters

- Anomalously shallow craters with fractured floors. [Schultz, 1976]
- 170 craters. [Jozwiak et al., 2012]
- Diameter range:
13 - 208 km. [Jozwiak et al., 2012]
- 8 morphologic subclasses.
[Schultz, 1976; Jozwiak et al., 2012]



Crater Humboldt, D= 207 km
(LOLA +LROC-WAC)

Floor-Fractured Crater Distribution, N=170

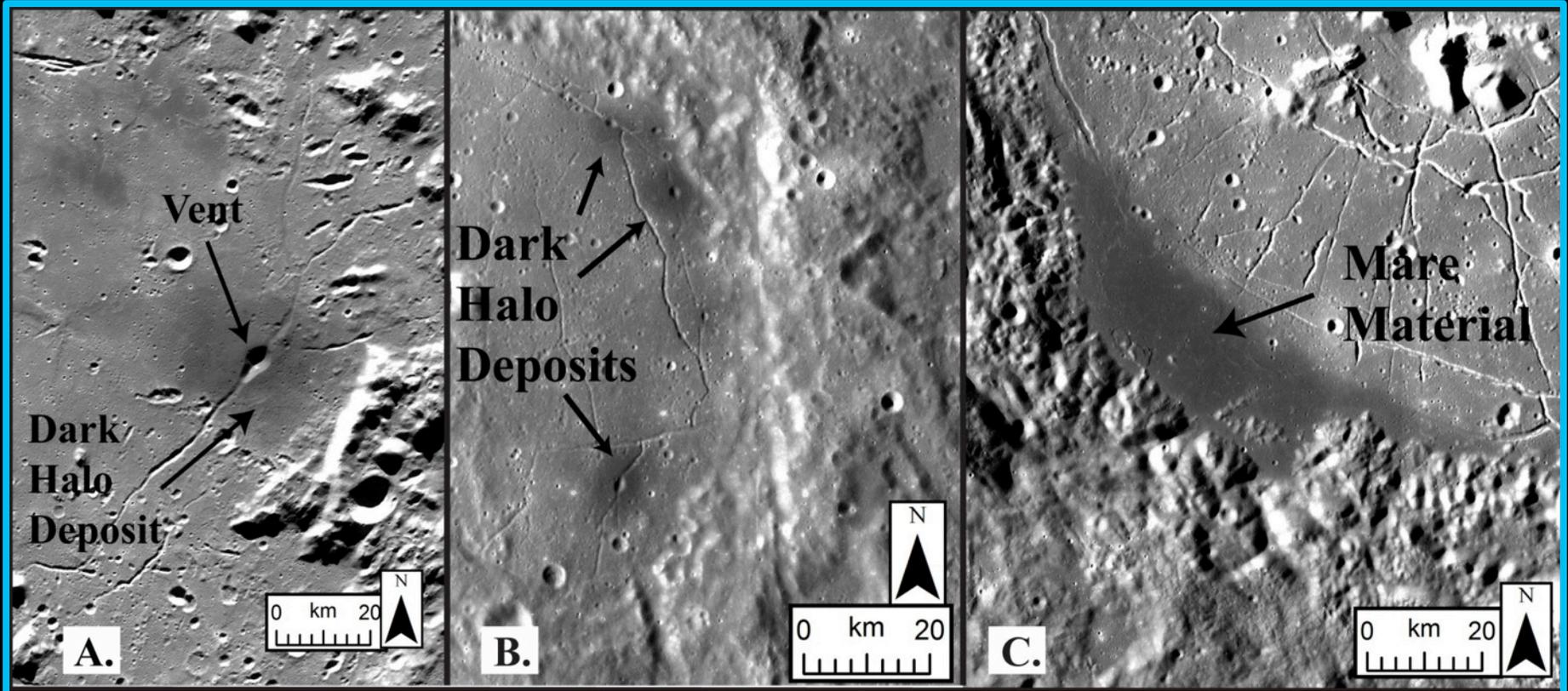


Floor-Fractured Crater Distribution: Jozwiak et al. (2012)

Mare Distribution: USGS

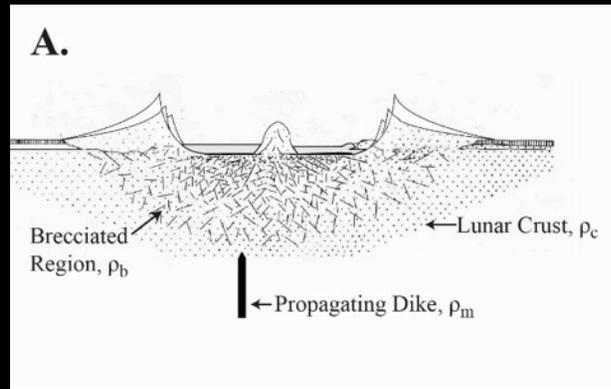
Cryptomare Distribution: Whitten and Head (2015)

Volcanic Morphologies

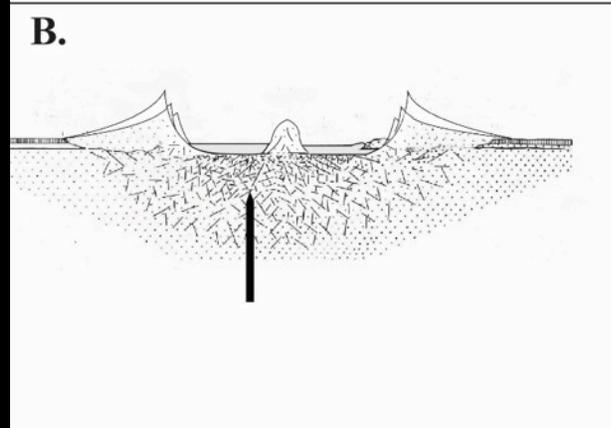


Magmatic Intrusion and Sill Formation

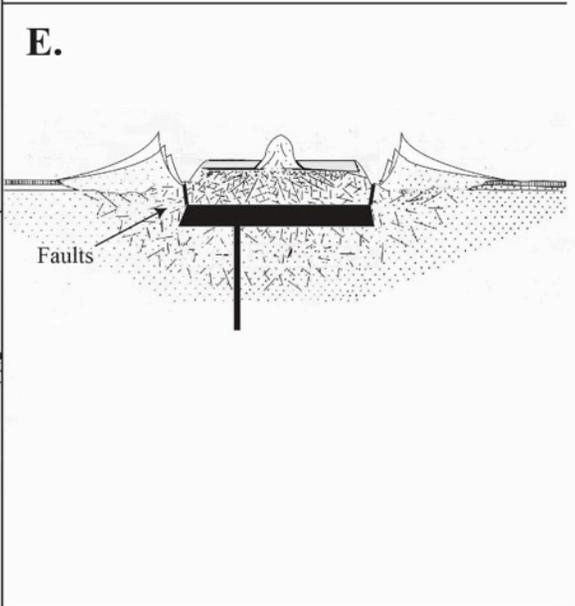
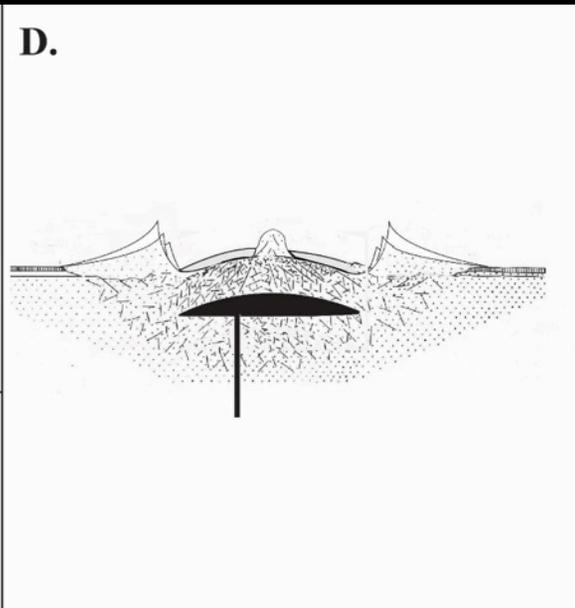
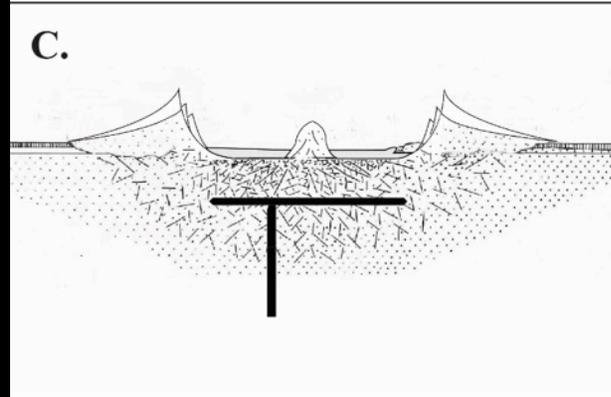
Dike
Propagation



Dike
Stalling



Sill
Formation



Domed
End-
Member
 $D < 40$ km

Flat End-
Member
 $D > 40$ km

Jozwiak et al.
(2015)

Not to scale.

Intrusive Magmatic Volume

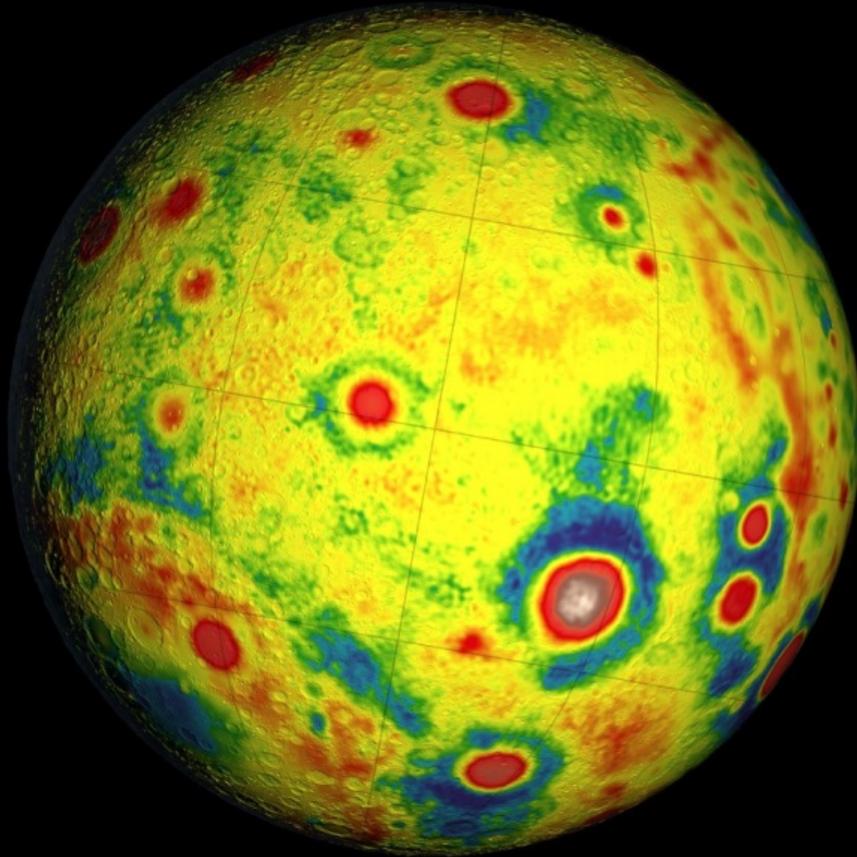
- Use floor morphology to estimate intrusion morphology—Domed vs. Tabular.
 - Supported by modeling of floor deformation [Thorey and Michaut, 2014].
- Intrusion thickness = $D_{theoretical} - D_{observed}$
 - Theoretical depth from Pike [1980].
 - Observed depth from LOLA.

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- Estimated volume $\sim 3 \cdot 10^6 \text{ km}^3$

Equivalent to 30% of the mare volume!

Investigating Intrusion Structure and Evolution

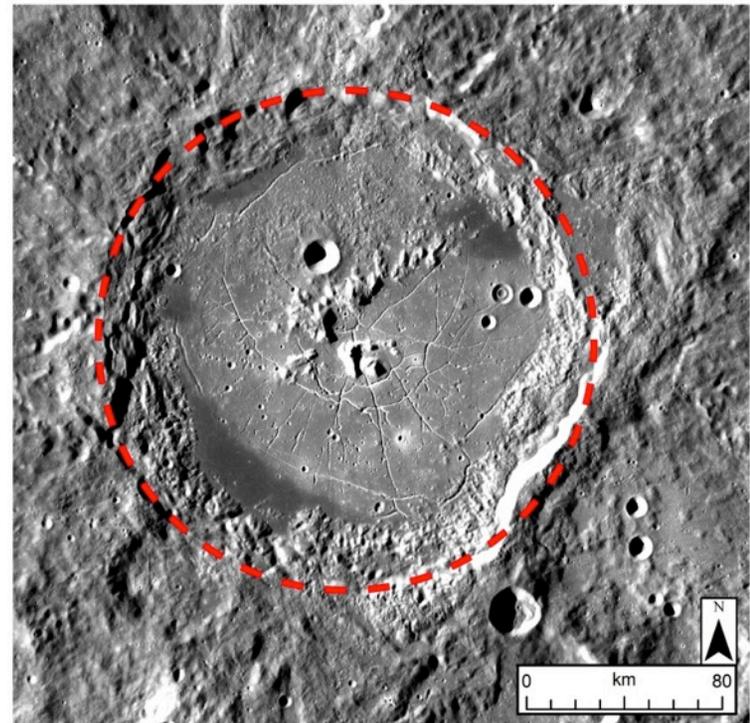
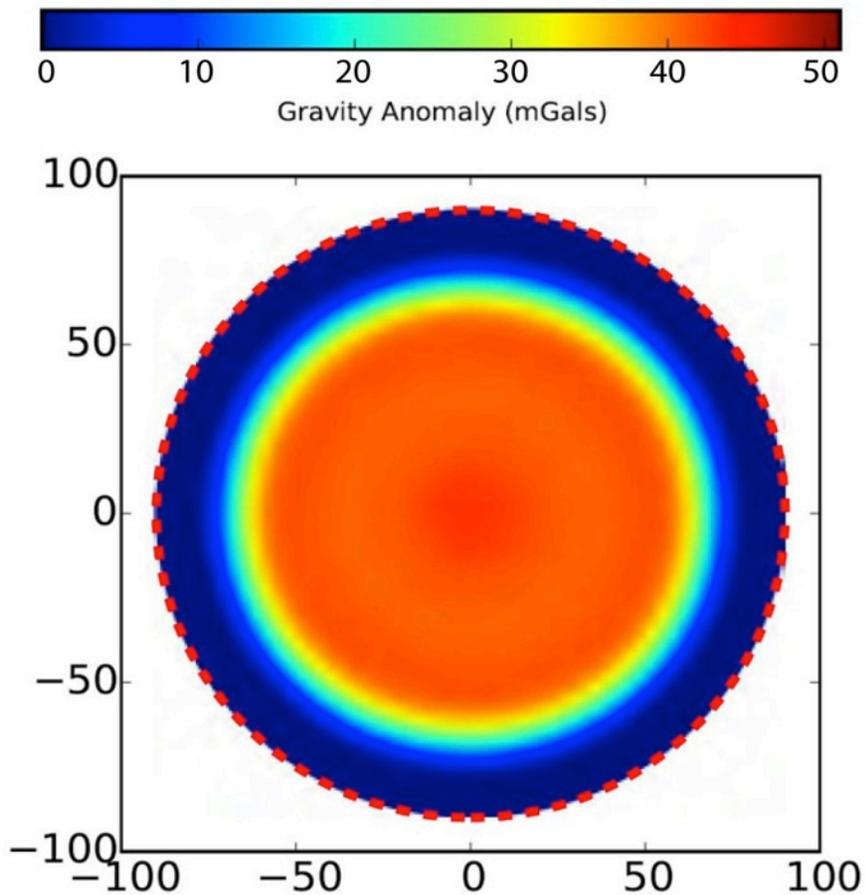


- Bouguer gravity data reveals variations in internal mass distribution.
- Mare basalt density $>$ lunar crustal density.
- Predict intrusions to have distinct, positive anomaly.

Predictions for Bouguer Anomaly

Humboldt

$D = 207$ km

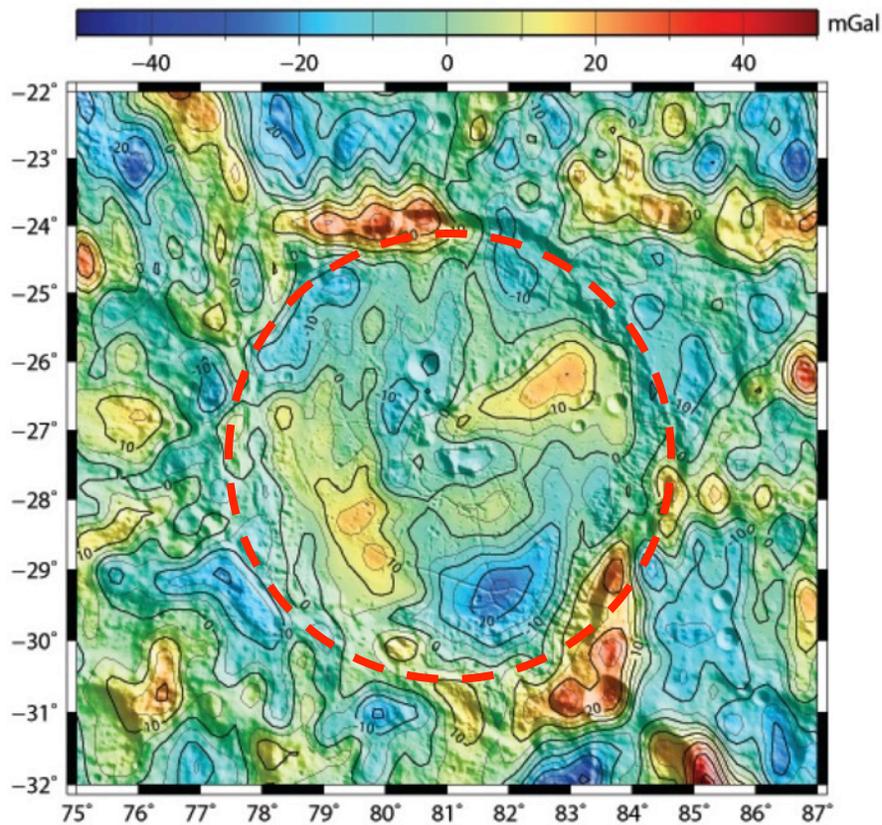


Synthetic Bouguer Gravity
Thorey et al. (2015)

Observed Bouguer Anomaly

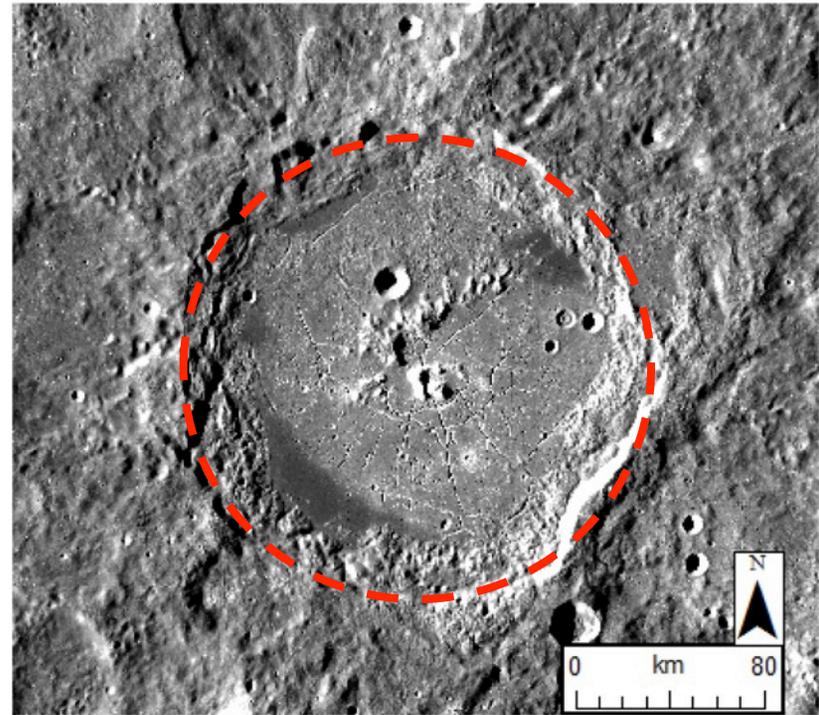
Humboldt

$D = 207 \text{ km}$



A.

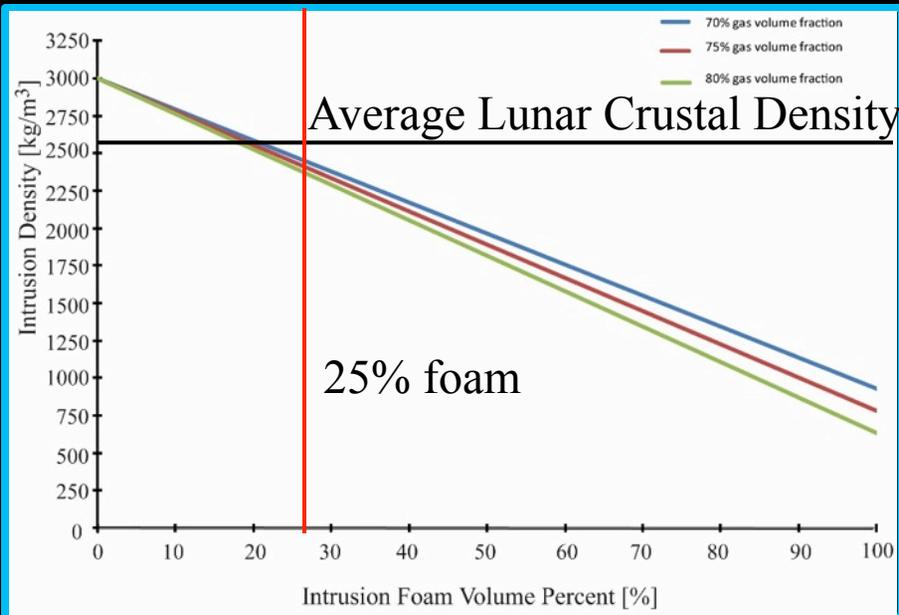
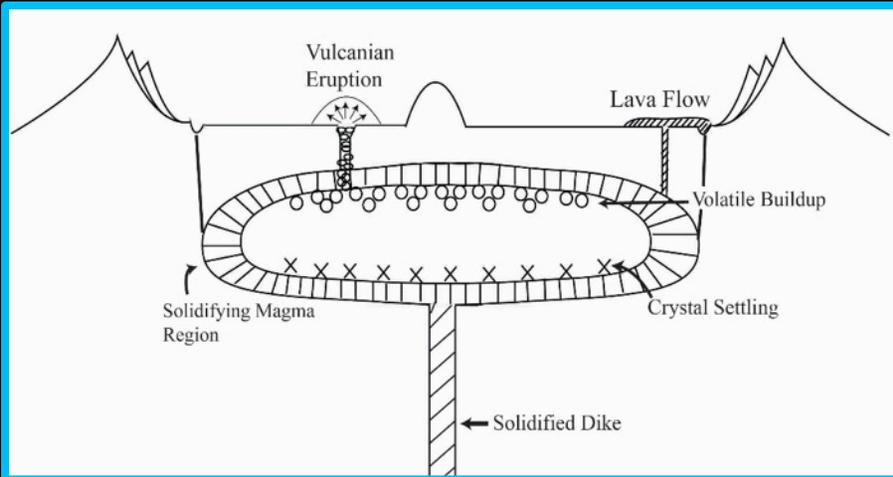
Orders 100-600



B.

LROC-WAC

Degassing of Intrusions

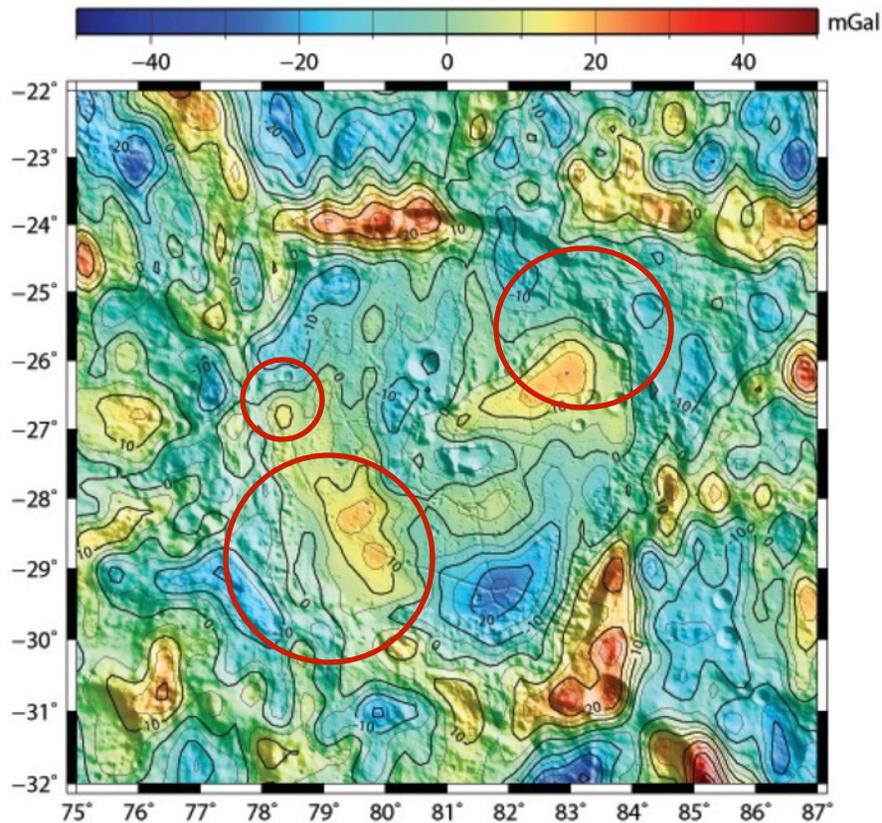


- Smelting reaction in magma between metal oxides and graphite forms CO at shallow depths, and dike propagation concentrates the bubbles in a foam at the dike tip. [Fogel and Rutherford, 1995]
- Additional volatile contributions from H₂O, Cl₂, and F₂ [Saal et al., 2008].
- **Bubbles form gas-rich foam at the intrusion roof.** [e.g. Jaupart and Vergnolle, 1989; Wilson and Head, 2003]

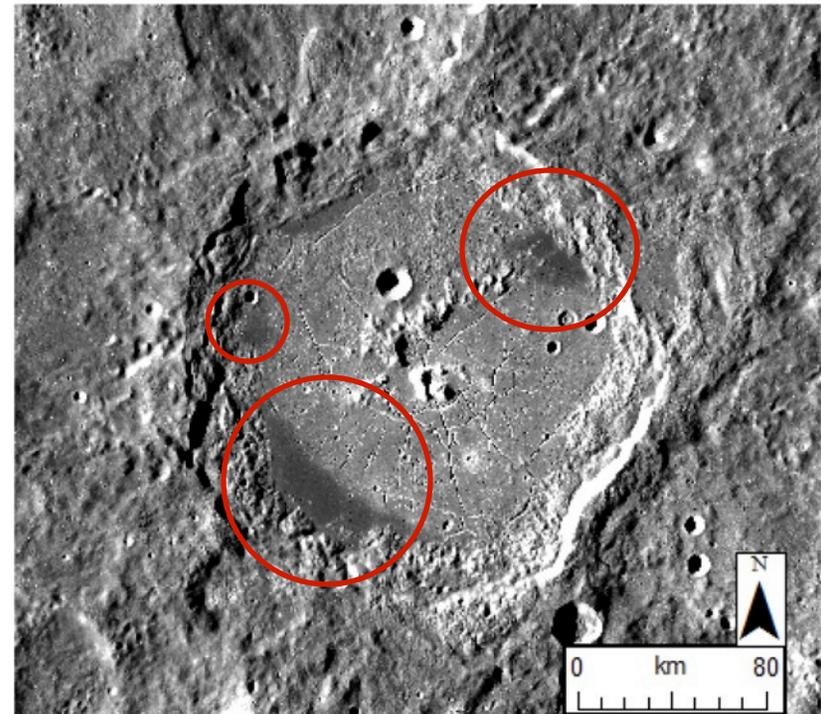
Correlation with Volcanic Features

Humboldt

$D = 207 \text{ km}$



A.



B.

Orders 100-600

LROC-WAC

Implications for Lunar Magmatic Intrusions

- Volatile buildup and degassing is a natural consequence of magmatic intrusion evolution. [Wilson and Head, 2003]
- Variations in foam gas density and foam volume have significant effect on overall intrusion density.
- Process explains observation that higher BA are correlated with surficial volcanic deposits, which would have been formed during initially explosive degassing of the local intrusion region.
- Passive venting can also result in foam loss and density changes.
- The heterogeneous nature of this process and the lunar crust in general means that magmatic intrusions suspected from gravity data can be better understood and often verified when studied in combination with geological data.